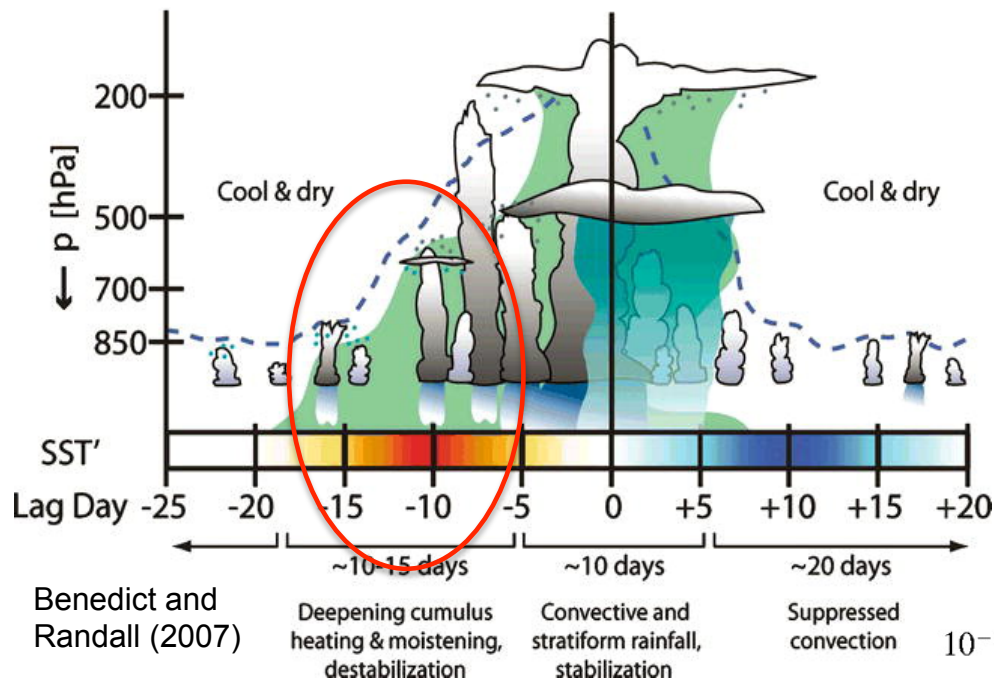


Convection Sensitivity to Atmospheric State During the Developing Stage of the MJO



**Tony Del Genio, Jingbo Wu, Yonghua Chen,
Daehyun Kim, Mao-Sung Yao**

NASA GISS and Columbia University



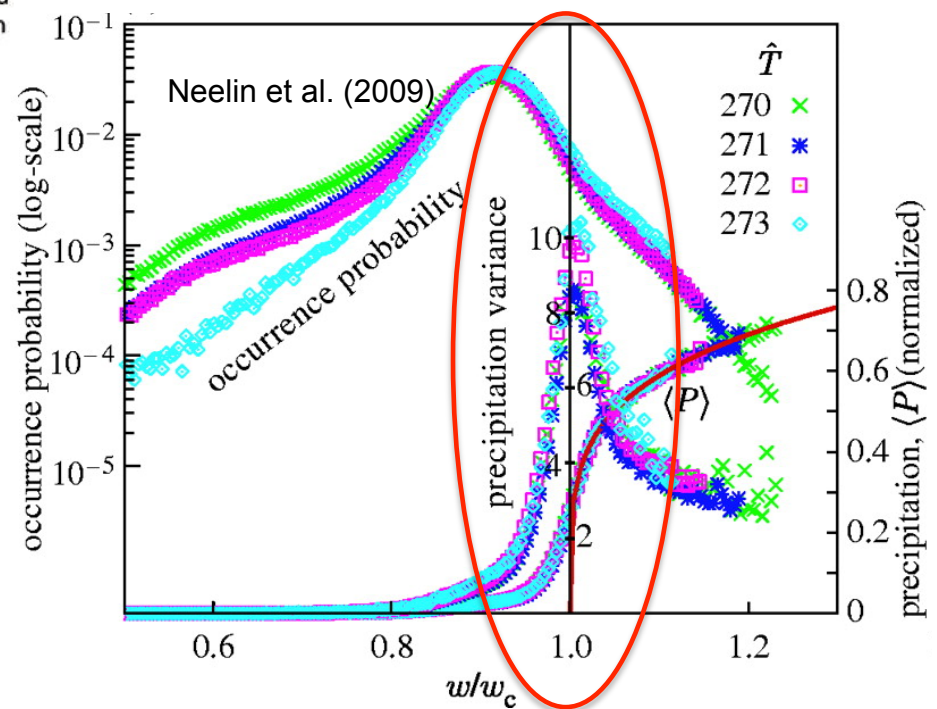
Emerging view of MJO initiation: moisture buildup by shallow/congestus clouds eventually triggering deep convection

A challenge for GCMs, which tend to simulate MJO poorly

But actually more interesting than simple picture suggests

TRMM TMI precipitation vs. column water vapor:

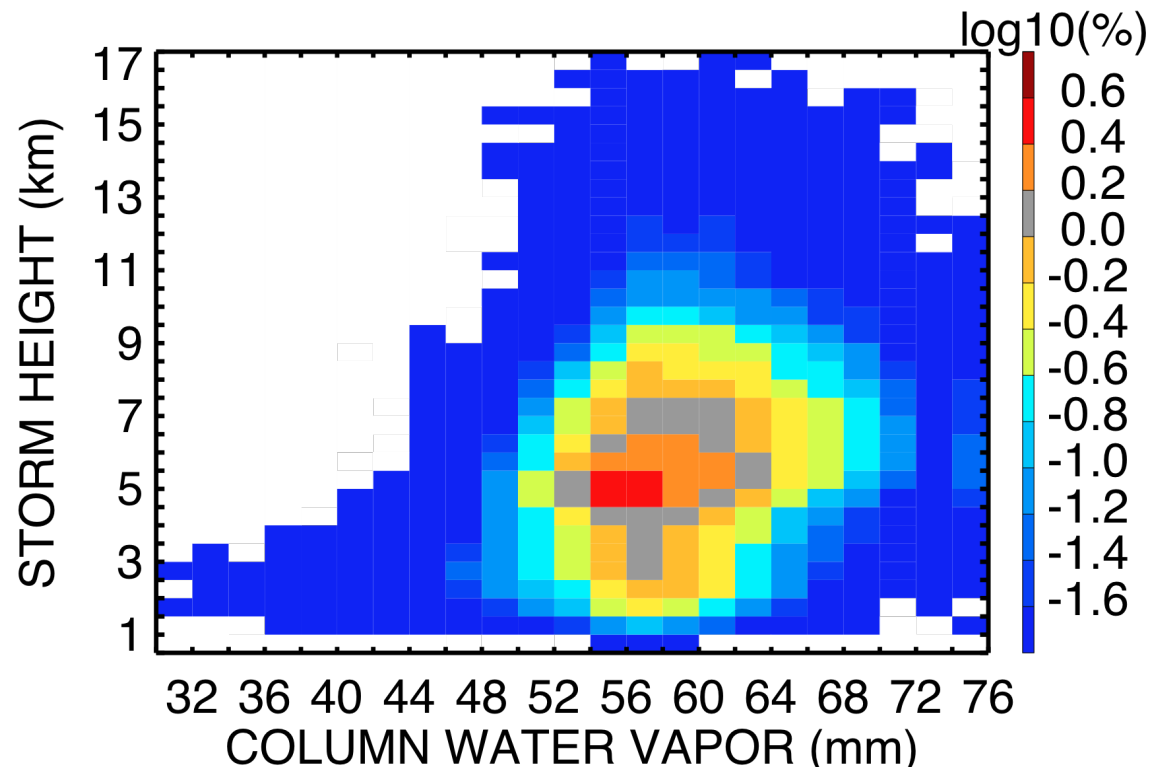
Sharp increase in P at "critical" CWV, maximum rainfall variance and occurrence probability near critical value



Approach

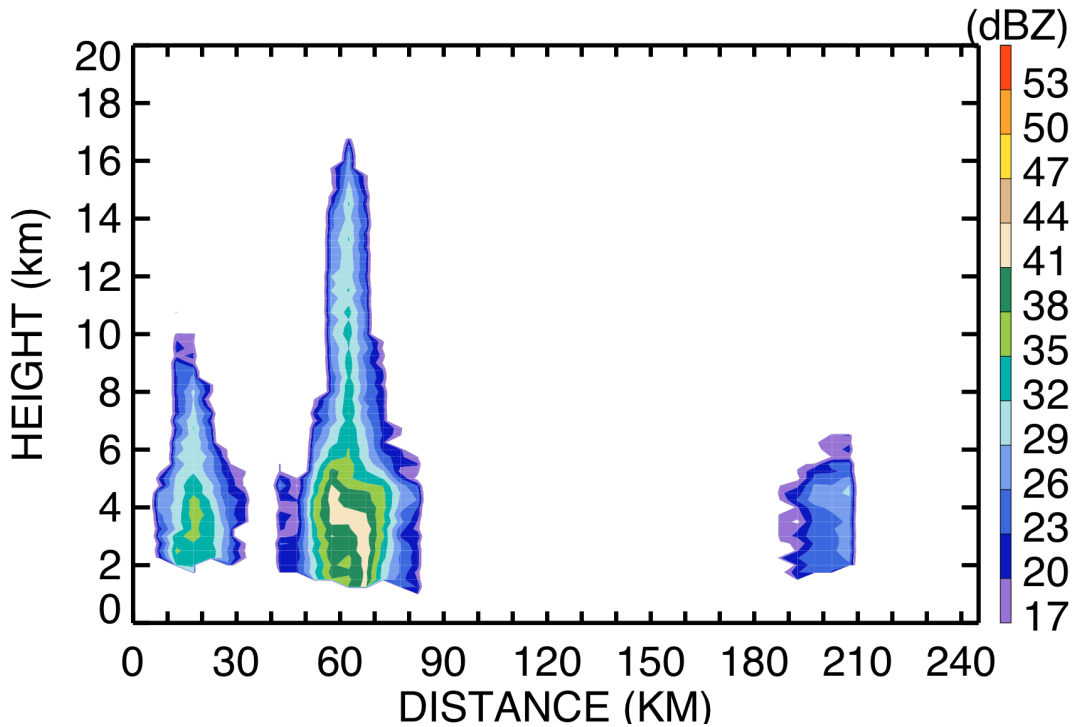
- **Wheeler-Hendon + NOAA CPC MJO indices: 10 boreal non-summer MJO peaks (2006-2010) in equatorial Indian Ocean, Maritime Continent, West Pacific; focus on transition from suppressed to disturbed phase (10-14 days before peak)**
- **TRMM PR storm heights + TMI column water vapor**
- **CloudSat/CALIPSO GEOPROF-LIDAR convective cloud top heights + AMSR-E column water vapor**
- **GISS Model E2 GCM CMIP5 version + stronger convective entrainment and rain evaporation**

PR storm height statistics during MJO transition phase



- Transition to deep Convection at CWV~ 50-66 mm
- Maximum convective depth variance in critical CWV range, explains P variance
- Also maximum frequency of occurrence near critical CWV

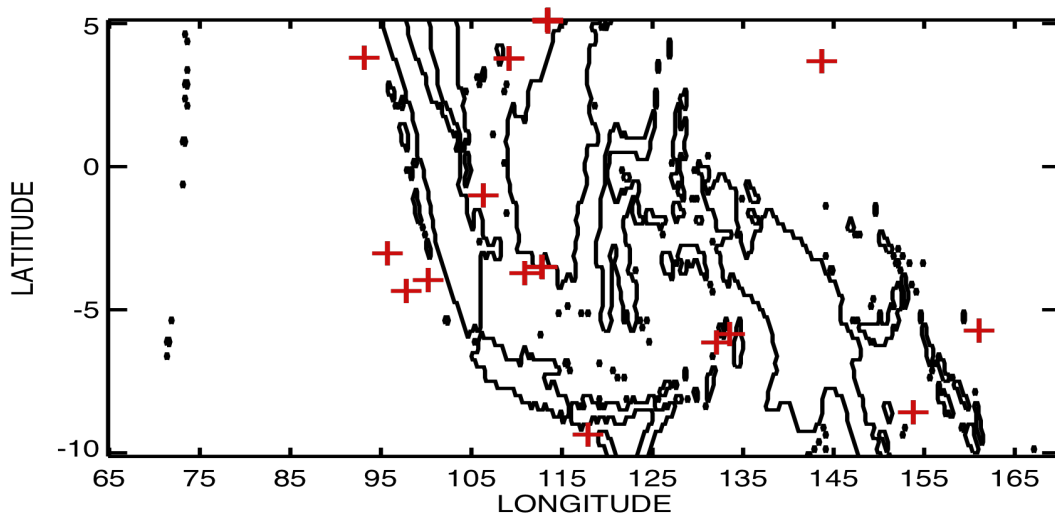
- Deepest storm heights at intermediate rather than wettest CWV – why?
- Some storms as deep as 16-17 km – are these real?



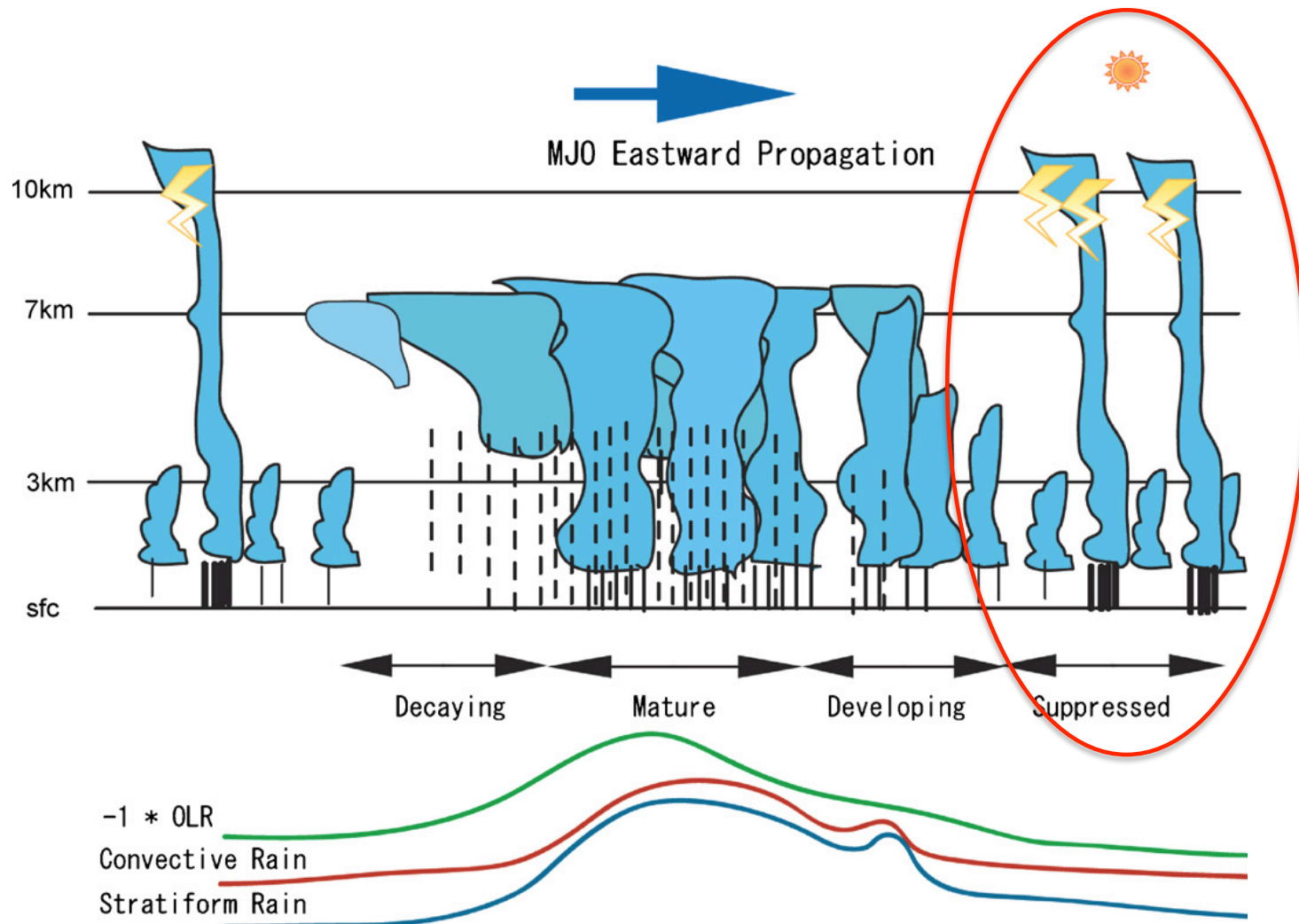
**Sample PR reflectivity profiles for one case:
~30 dBZ up to 15 km**

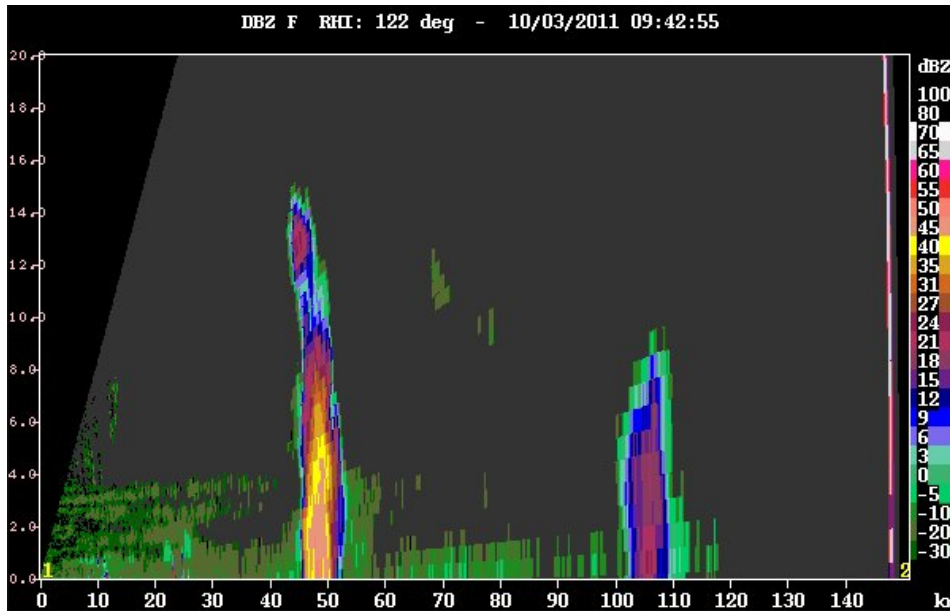
Isolated cells

All located in Maritime Continent, over very warm water, maybe continental influence?

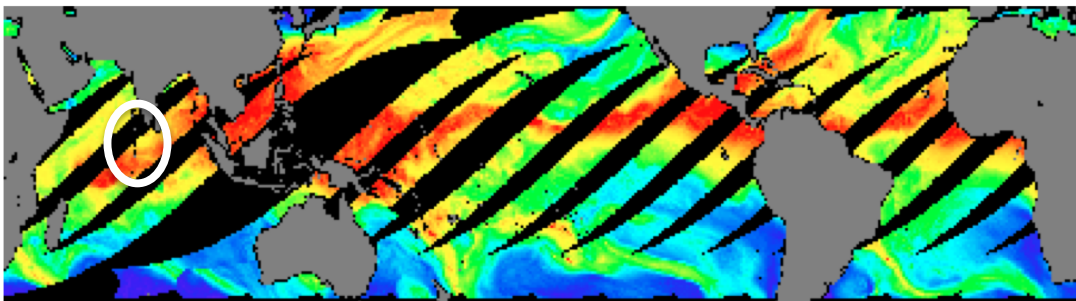


Morita et al. (2006): ~11 km TRMM PR rain top heights and LIS lightning during suppressed MJO phase

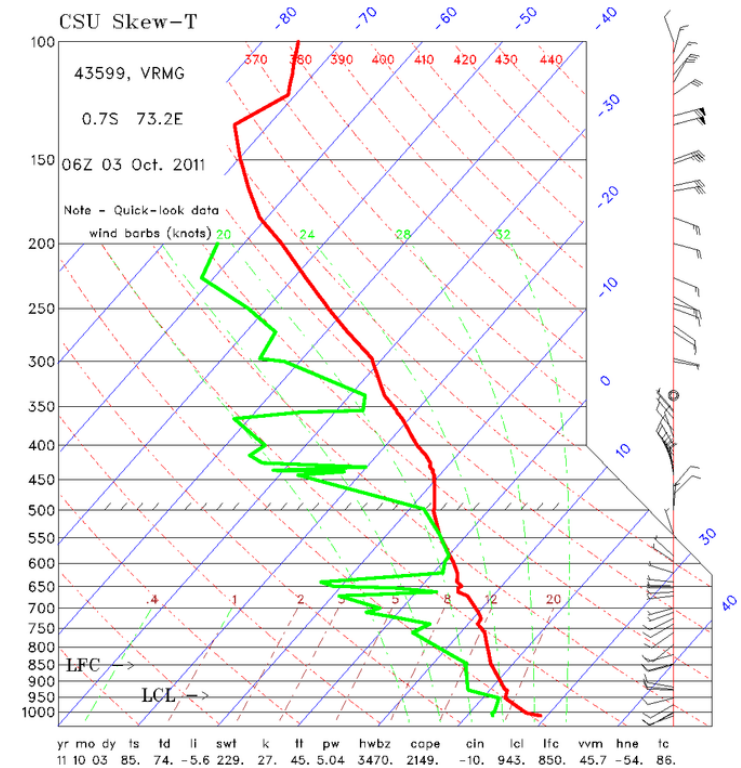




TMI CWV 10/3/11



***Thanks to Eric Maloney and Bob Houze for pointing out this case on the DYNAMO blog**

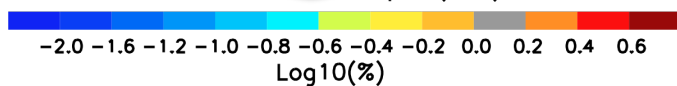
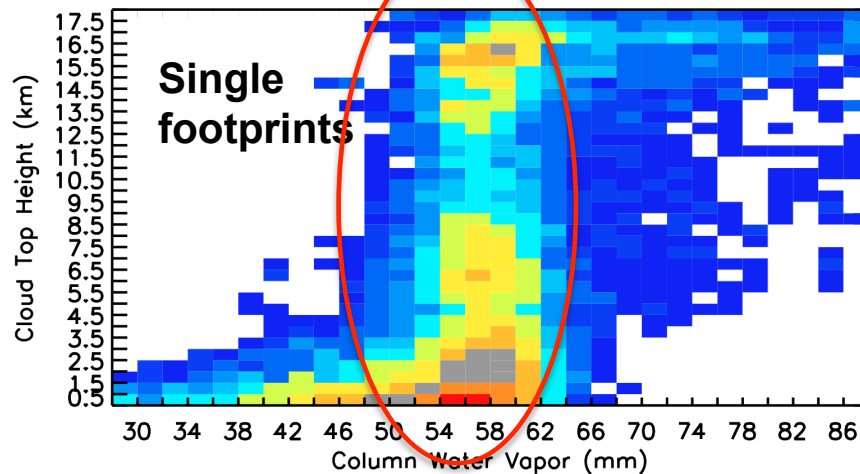
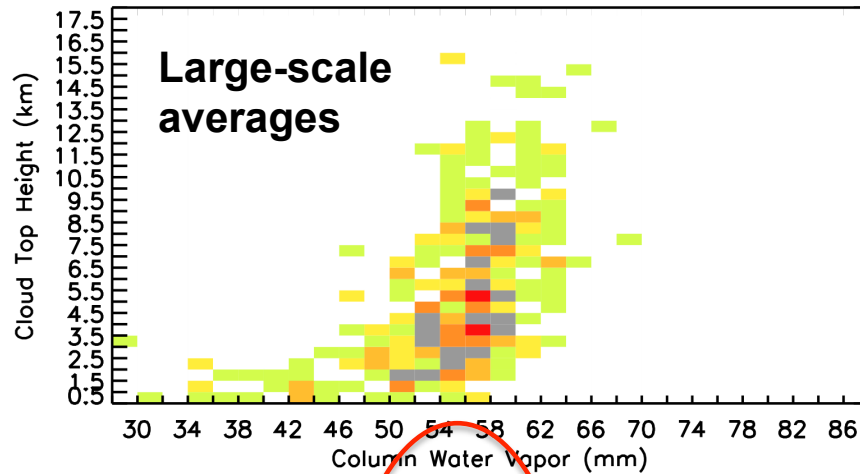


DYNAMO Gan Island case, 10/3/11*:

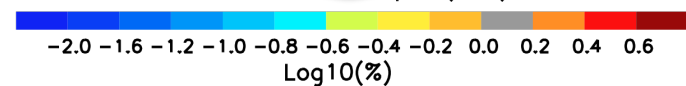
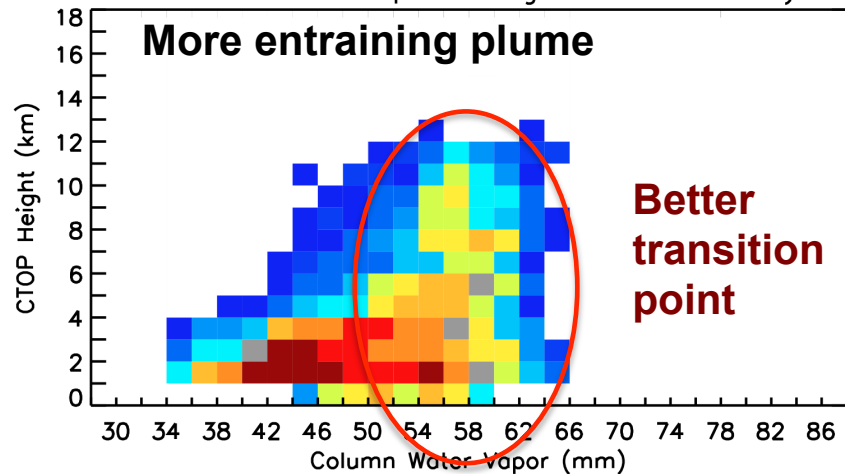
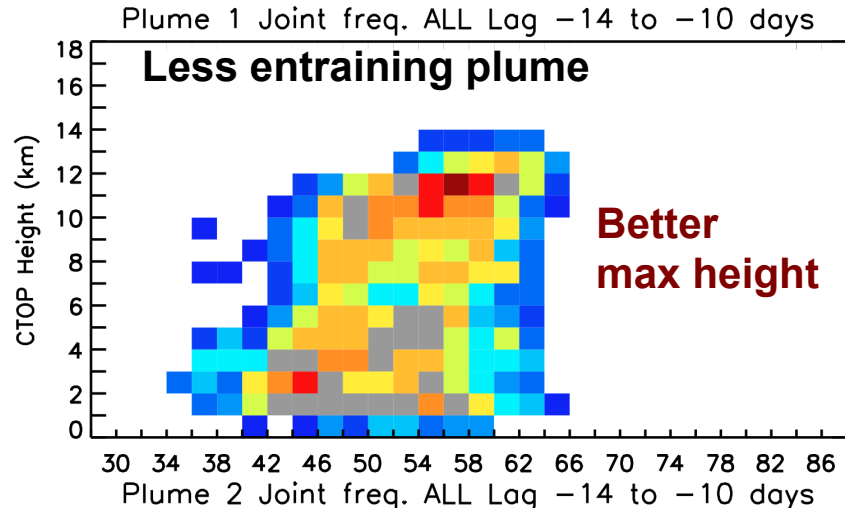
- Fairly low CWV due to mid-troposphere dry layers
- But fairly unstable lapse rate

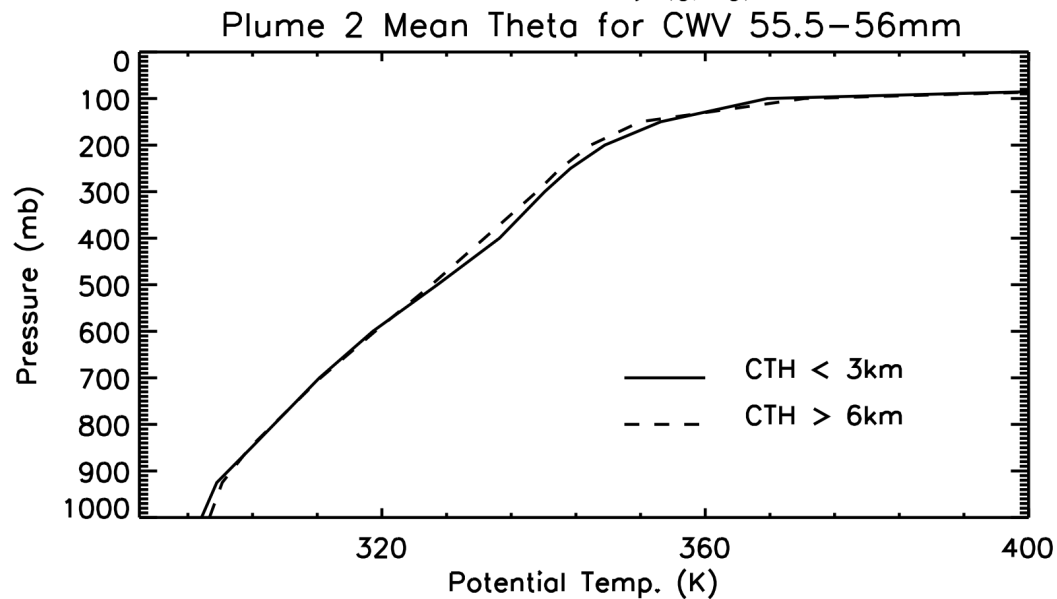
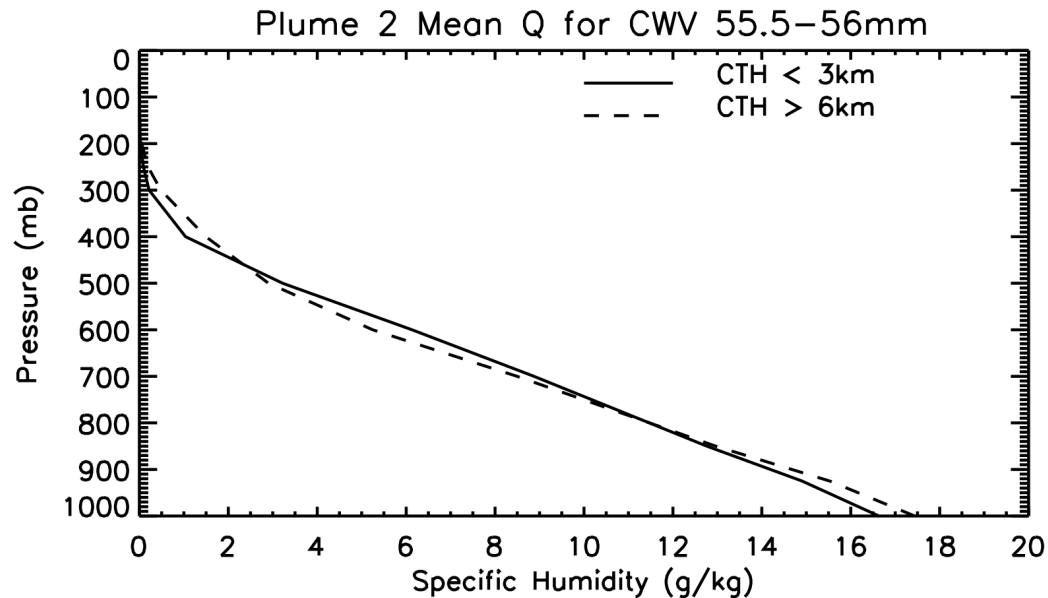
CloudSat/CALIPSO convective cloud top heights: Deepest tops at wettest CWV just as deep as those at lower CWV; GCM can match data but only with strong entrainment

Data



GCM





How does GCM produce convection depth variance for a given value of CWV with a deterministic parameterization?

Deeper when PBL is a bit warmer/wetter

Gregory entrainment scheme ($\epsilon \sim B/w^2$) implies:

**Warmer/wetter rising parcel
-> stronger updraft ->
weaker entrainment ->
better able to penetrate
drier mid-troposphere**

Summary

- **TRMM PR storm heights indicate large variability in convection depth at intermediate CWV during transition from suppressed to disturbed phase of MJO**
- **Storm heights deeper when column is moderately moist than when it is most humid; associated with isolated convective cells**
- **Comparison to CloudSat/CALIPSO indicates that peak convection depth is similar at intermediate and high CWV; PR storm height behavior thus suggests stronger storms in more unstable environment when CWV is lower**
- **GISS GCM is capable of reproducing observed transition from shallow to deep convection and associates deeper convection with warmer, wetter PBL; but weak entrainment must be restricted to very specific conditions**