

Basics

$$N(D) = \frac{N_t}{D_c} g(x) \quad \text{with} \quad x = \frac{D}{D_c}$$

$$g(x; \lambda, \mu) = \frac{\lambda^{\mu+1}}{\Gamma(\mu+1)} x^\mu \exp(-\lambda x)$$

$$N(D) = \frac{N_t}{D_c} g(x; \lambda, \mu) = \frac{N_t}{D_c} \left[\frac{\lambda^{\mu+1}}{\Gamma(\mu+1)} \left(\frac{D}{D_c} \right)^\mu \exp\left(-\lambda \frac{D}{D_c}\right) \right]$$

$$D_c = D_{4,3} = \frac{M_4}{M_3}$$

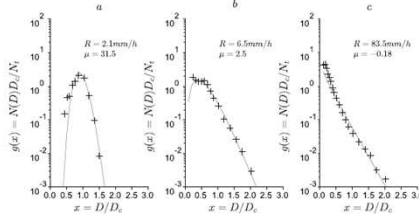
$$M_k = \int_0^\infty N(D) D^k dD = \frac{\Gamma(\mu+k+1) N_t D_c^k}{\Gamma(\mu+1) \lambda^k}$$

$$\lambda = \mu + 4,$$

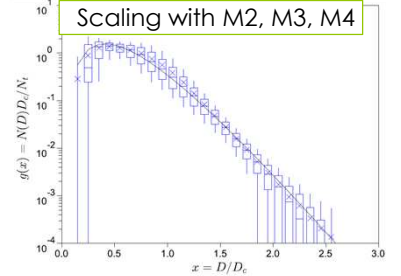
$$M_0^* = \frac{M_2 \lambda^2}{(\mu+2)(\mu+1) D_c^2}$$

$$\mu = \frac{3-4A}{A-1} \quad \text{with} \quad A = \frac{M_3^2}{M_2 M_4}$$

Fits for individual spectra ...



and for a 3-year DSD climatology



Comments

This study offers a unified formulation of single and double-moment normalizations of the raindrop size distribution (DSD), which have been proposed in the framework of the scaling analysis in the literature.

The key point is to consider a well-defined "general distribution" $g(x)$ as the probability density function (pdf) of the raindrop diameter scaled by a characteristic diameter D_c . We use the ratio of the 4th to the 3rd DSD moments as the characteristic diameter and the two-parameter gamma pdf to model the $g(x)$ -function.

The theory is illustrated with a three-year DSD time series collected with a Parsivel disdrometer, including a large variety of convective and non-convective rainfall events representative of the rainfall climatology in the Cevennes region, France. It is first shown that three DSD moments (M_2, M_3, M_4) allow to satisfactory model the DSD both for individual spectra and for time series of spectra.

The formulation is then extended to the one- and two-moment normalization by introducing single and dual power-law models between the explained moments (total concentration, characteristic diameter) and the scaling moment(s). Compared with previous scaling formulations, our approach explicitly accounts for the prefactors of the power-law models to yield a unique and dimensionless $g(x)$, whatever the scaling moment(s) considered. A parameter estimation procedure, based on the analysis of the power-law regressions and the so-called self-consistency relationships, is proposed for both the one- and two-moment normalizations.

When implemented with contrasted scaling DSD moments (rainrate and/or radar reflectivity), the method yields $g(x)$ -functions consistent with the one obtained with the three-moment normalization, although the 3-year DSD time series exhibits quite a large variability.

The intra-event variability of the DSD is illustrated for 22 October 2008 rain event: it is shown that very consistent $g(x)$ -functions can be obtained for homogeneous rain phases, whatever the scaling moments used, and that the $g(x)$ -functions may present contrasting shapes from one phase to another. This supports the idea that the $g(x)$ -function is process-dependent and not unique as hypothesized in the scaling theory.

Single-moment normalization

$$N_t = C_i M_i^{\alpha_i}$$

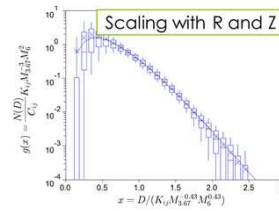
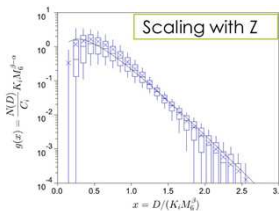
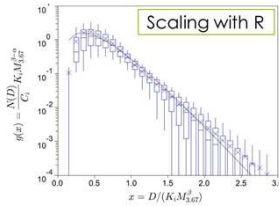
$$D_c = K_i M_i^{\beta_i}$$

$$N(D) = \frac{C_i M_i^{\alpha_i}}{K_i M_i^{\beta_i}} \frac{\lambda^{\mu+1}}{\Gamma(\mu+1)} \left(\frac{D}{K_i M_i^{\beta_i}} \right)^\mu \exp\left(-\lambda \frac{D}{K_i M_i^{\beta_i}}\right)$$

$$M_k = \frac{\Gamma(\mu+k+1)}{\Gamma(\mu+1)} C_i K_i^k \frac{M_i^{\alpha_i+k\beta_i}}{\lambda^k}$$

$$\alpha_i + i\beta_i = 1,$$

$$\frac{\Gamma(\mu+i+1)}{\Gamma(\mu+1)} C_i \left(\frac{K_i}{\lambda} \right)^i = 1.$$



Double-moment normalization

$$N_t = C_{ij} M_i^{\alpha_i} M_j^{\alpha_j}$$

$$D_c = K_{ij} M_i^{\beta_i} M_j^{\beta_j}$$

$$N(D) = \frac{C_{ij} M_i^{\alpha_i} M_j^{\alpha_j}}{K_{ij} M_i^{\beta_i} M_j^{\beta_j}} \frac{\lambda^{\mu+1}}{\Gamma(\mu+1)} \left(\frac{D}{K_{ij} M_i^{\beta_i} M_j^{\beta_j}} \right)^\mu \exp\left(-\lambda \frac{D}{K_{ij} M_i^{\beta_i} M_j^{\beta_j}}\right)$$

$$M_k = \frac{\Gamma(\mu+k+1)}{\Gamma(\mu+1)} C_{ij} K_{ij}^k \frac{M_i^{\alpha_i+k\beta_i} M_j^{\alpha_j+k\beta_j}}{\lambda^k}$$

$$\alpha_i + i\beta_i = 1,$$

$$\alpha_j + i\beta_j = 0,$$

$$\alpha_i + j\beta_j = 0,$$

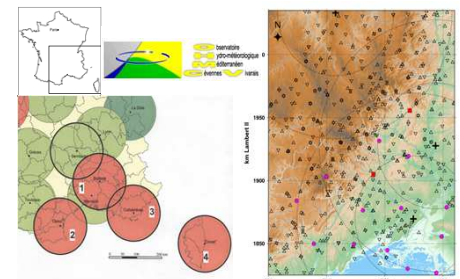
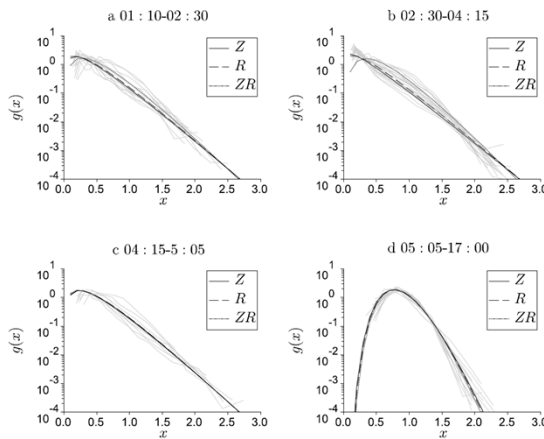
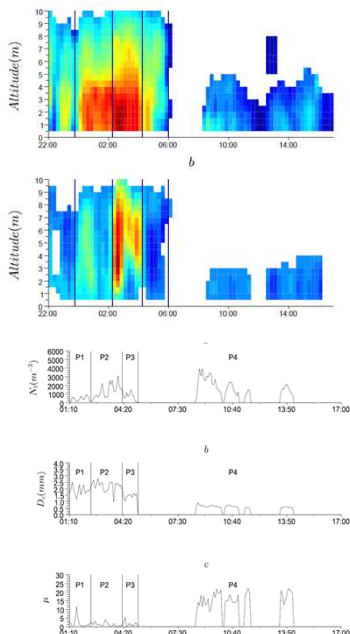
$$\alpha_j + j\beta_j = 1,$$

$$\frac{\Gamma(\mu+i+1)}{\Gamma(\mu+1)} C_{ij} \left(\frac{K_{ij}}{\lambda} \right)^i = 1,$$

$$\frac{\Gamma(\mu+j+1)}{\Gamma(\mu+1)} C_{ij} \left(\frac{K_{ij}}{\lambda} \right)^j = 1.$$

Normalization framework		μ
N_t and D_c		2.76
Z		1.47
R		2.00
ZR		2.80

DSD intra-event variability: the 22 October 2008 event



References:

1. Lee, G., I. Zawadzki, W. Szymer, D. Sempere Torres, and R. Uijlenhoet, 2004: A general approach to double-moment normalization of drop size distributions. *J. Appl. Meteorol.*, **43** (2), 364-381.
 2. Sempere Torres, D., J. M. Porrà, and J.-D. Creutin, 1994: A general formulation for raindrop size distribution. *J. Appl. Meteorol.*, **33** (12), 1494-1502.
 3. Sempere Torres, D., J. M. Porrà, and J.-D. Creutin, 1998: Experimental evidence of a general description for raindrop size distribution properties. *J. Geophys. Res.*, **103** (D2), 1785-1797.
 4. Yu, N., B. Boudevillain, G. Delrieu, and R. Uijlenhoet, 2012: Estimation of rain kinetic energy from radar reflectivity factor and/or rain rate based on a scaling formulation of the rain drop size distribution. *Water Resour. Res.*, **48**, W04505.