

# Ice-Phase Particle Size Distributions in May 20 MC3E Convective System

## Sensitivity to Ice Collection Kernel

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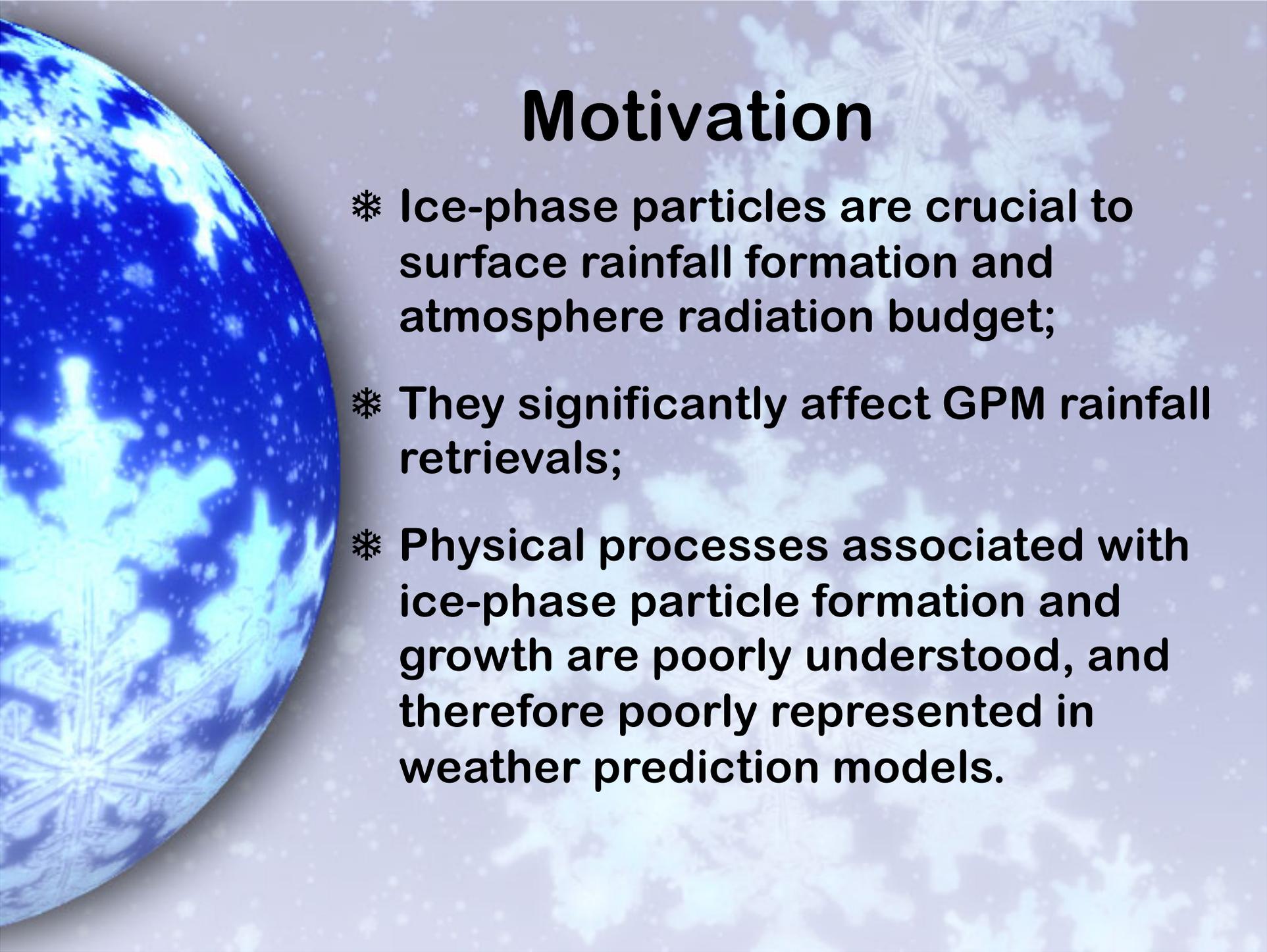
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<sup>3</sup>University of Maryland, College Park

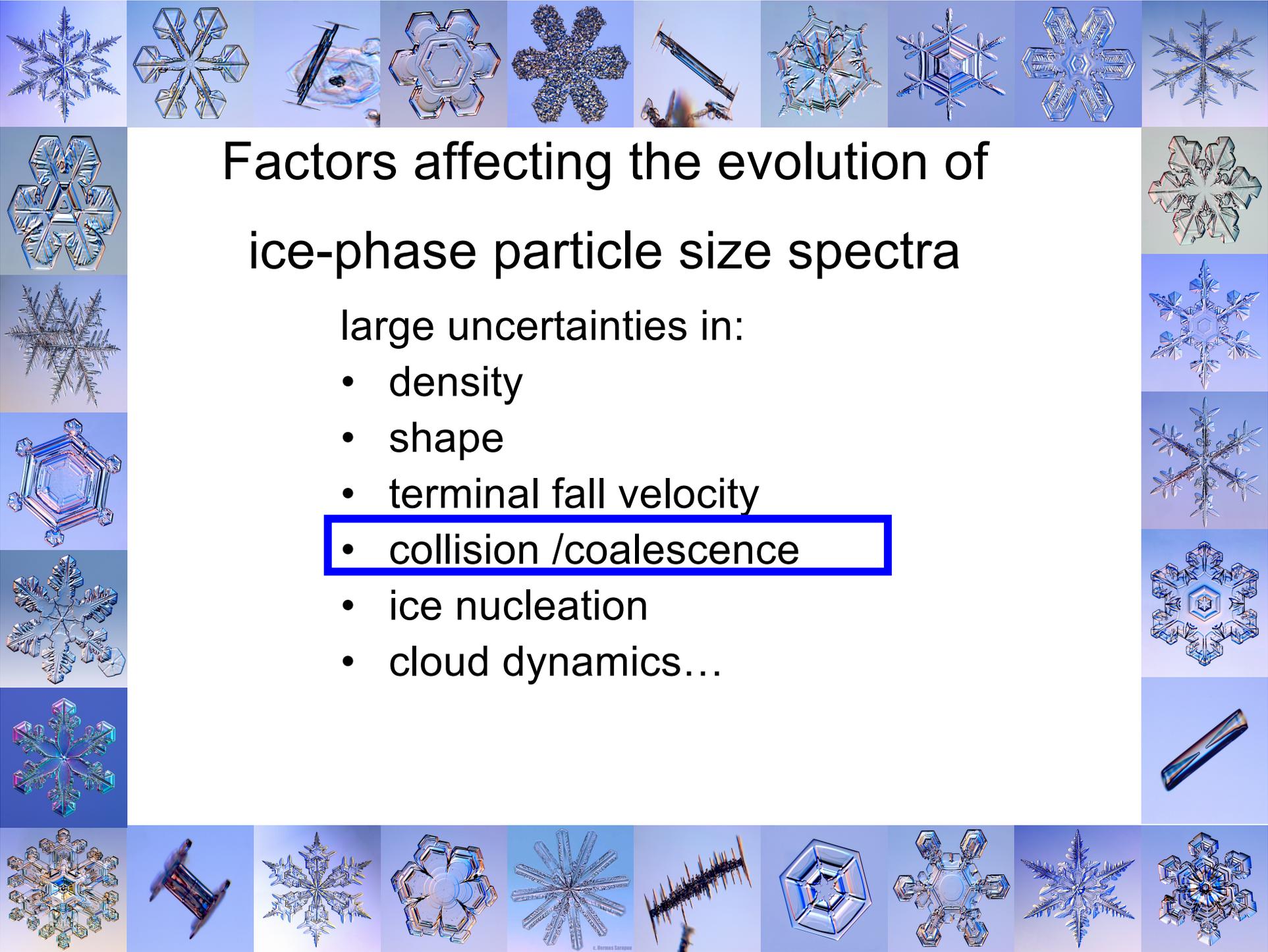
<sup>4</sup>Science System and Application, Inc

<sup>5</sup>The University of Arizona



# Motivation

- ❄ Ice-phase particles are crucial to surface rainfall formation and atmosphere radiation budget;
- ❄ They significantly affect GPM rainfall retrievals;
- ❄ Physical processes associated with ice-phase particle formation and growth are poorly understood, and therefore poorly represented in weather prediction models.



# Factors affecting the evolution of ice-phase particle size spectra

large uncertainties in:

- density
- shape
- terminal fall velocity
- collision /coalescence
- ice nucleation
- cloud dynamics...

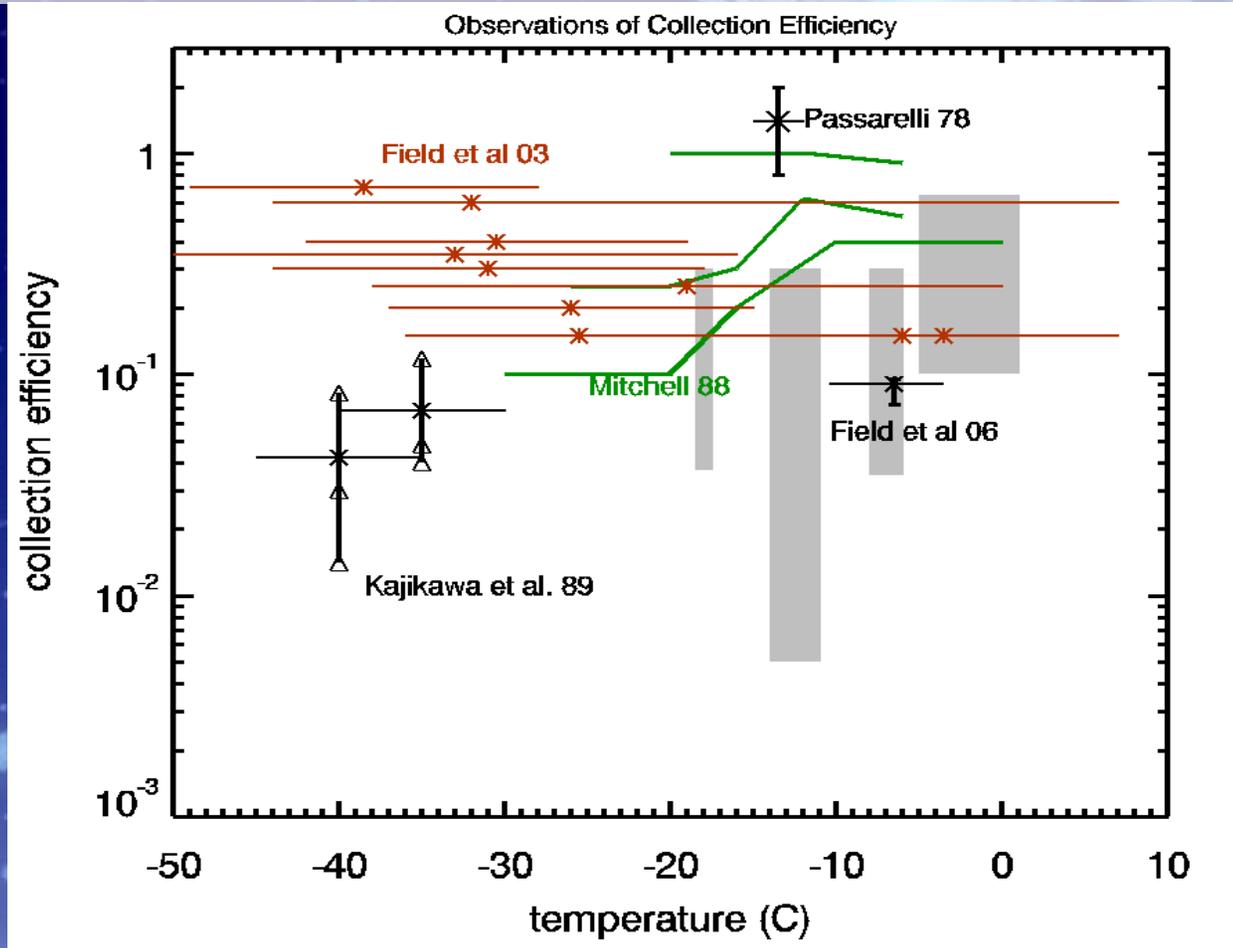
Collision/Coalescence process is governed by:

$$\frac{dn_k}{dt} = \frac{1}{2} \int_0^{m_k = m_i + m_j} K_{k-j,j} n_{k-j} n_j dm_j - n_k \int_0^{\infty} K_{kj} n_j dm_j$$

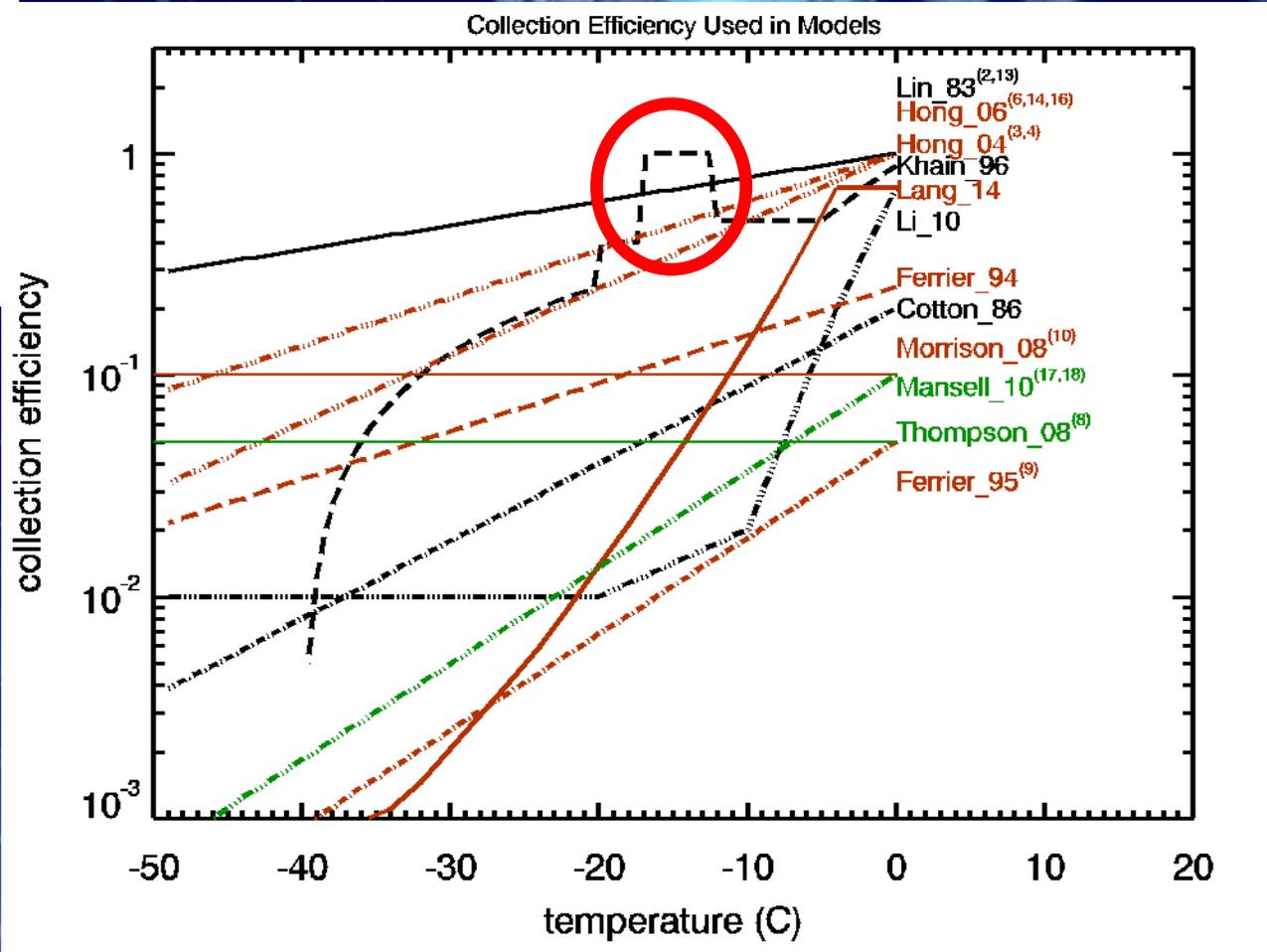
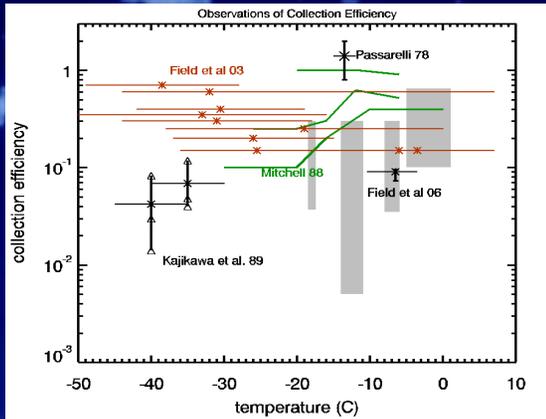
The collection kernel between two particles:

$$K_{ij} = \frac{\pi}{4} \underbrace{(D_i + D_j)^2}_{\text{size}} \underbrace{|v_i - v_j|}_{\text{Fall velocity}} \underbrace{E_c}_{\text{Collection Efficiency (T, RH)}}$$

# Collection Efficiencies $E_c(T)$ in Observations



# Collection Efficiency in Models



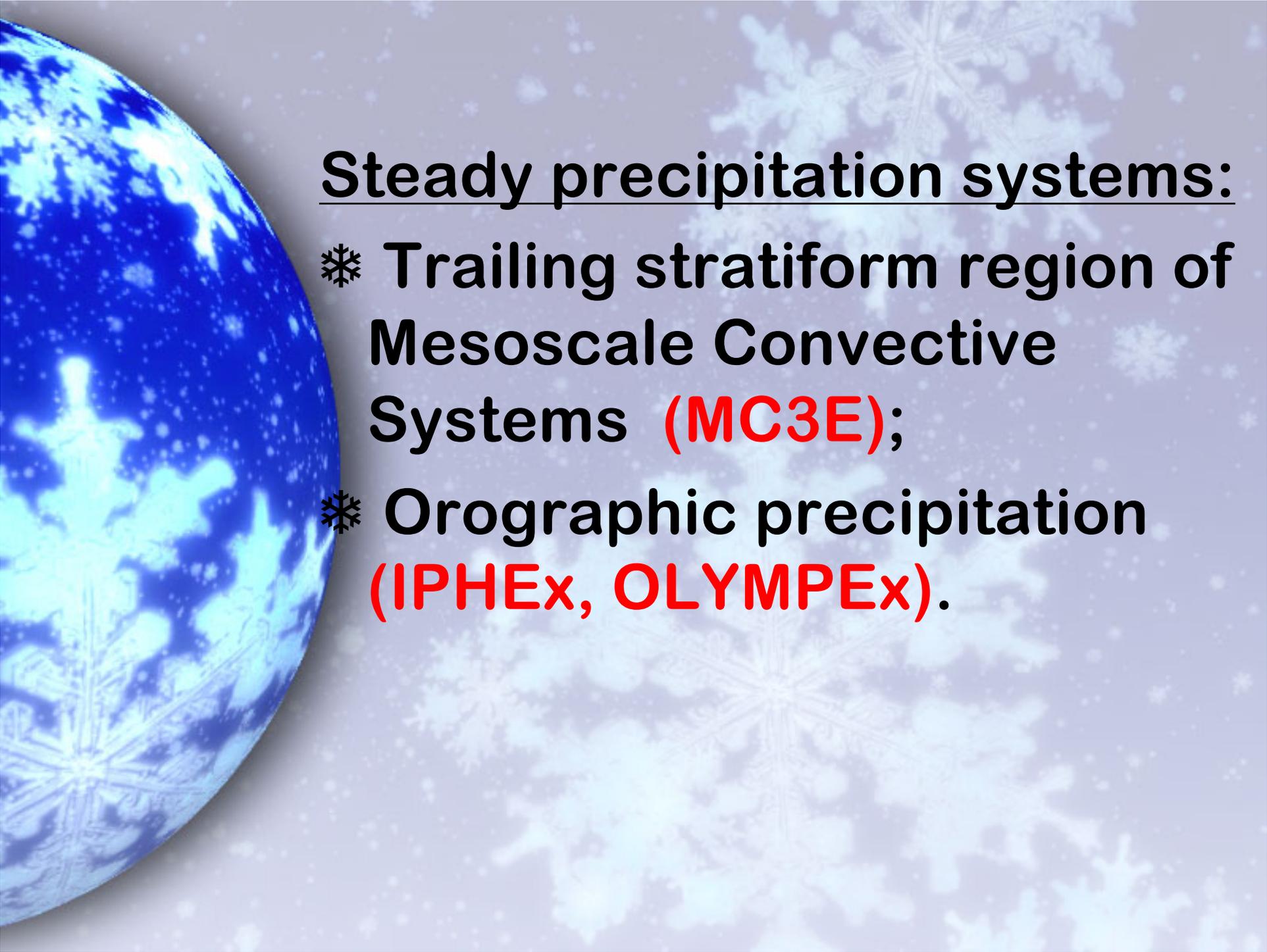


# Study Goal

**To Reduce Uncertainties in  $E_c(T)$**

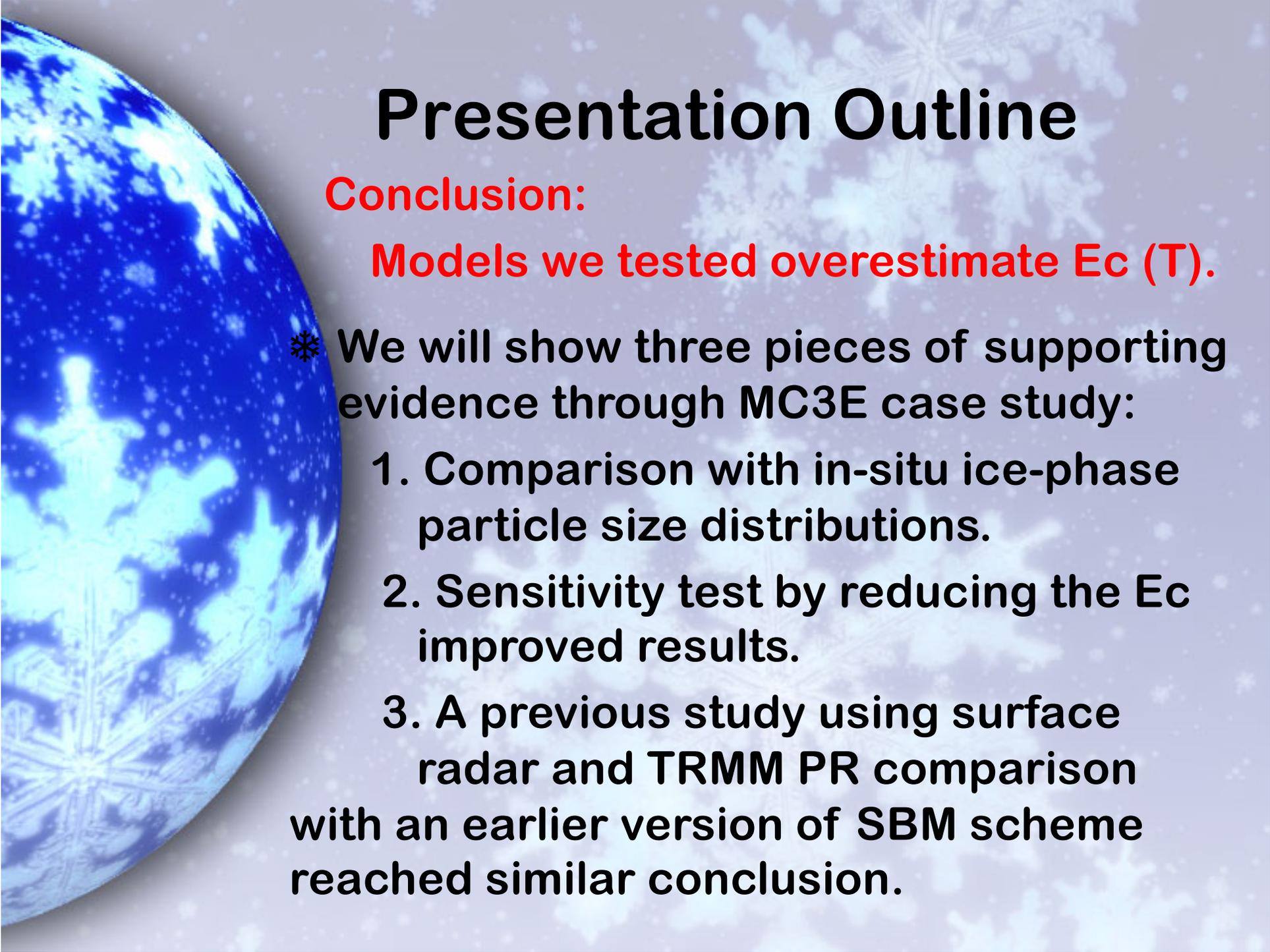
Using case studies with:

- ❄ In-situ PSD observations;
- ❄ Radar and GPM profiles;
- ❄ Cloud-resolving model with a detailed spectral bin microphysics scheme.



## Steady precipitation systems:

- ❄ Trailing stratiform region of Mesoscale Convective Systems (**MC3E**);
- ❄ Orographic precipitation (**IPHEX, OLYMPEX**).



# Presentation Outline

## Conclusion:

**Models we tested overestimate  $E_c$  (T).**

❄ We will show three pieces of supporting evidence through MC3E case study:

1. Comparison with in-situ ice-phase particle size distributions.
2. Sensitivity test by reducing the  $E_c$  improved results.
3. A previous study using surface radar and TRMM PR comparison with an earlier version of SBM scheme reached similar conclusion.

# Model Description

- WRF v3.6.1, single domain, 24-hour simulation;
- 2-km horizontal resolution, 44 vertical levels;
- Initiated with NCEP FNL Operational Global Analysis starting at 00 UTC, May 20, 2011;

- Hebrew U Spectral Bin Microphysical Scheme (SBM):

43 mass-doubling size bins representing particle size distributions for 3 shapes of ice crystals (column, plate, dendrite), snow aggregate, graupel, hail and cloud/rain.

- Morrison 2-Moment Microphysical Scheme:

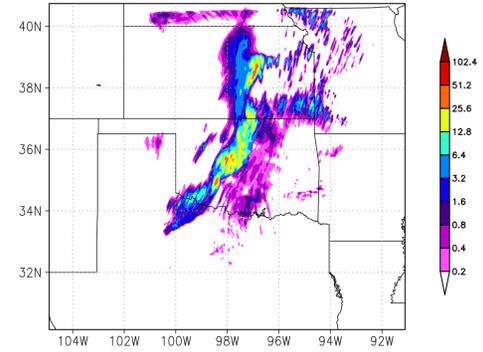
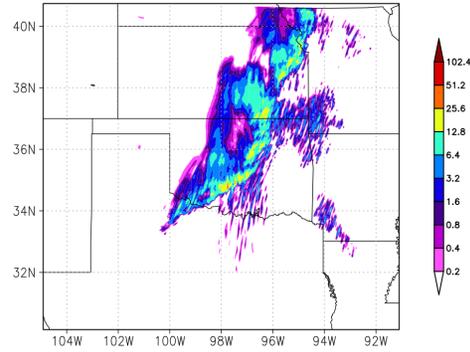
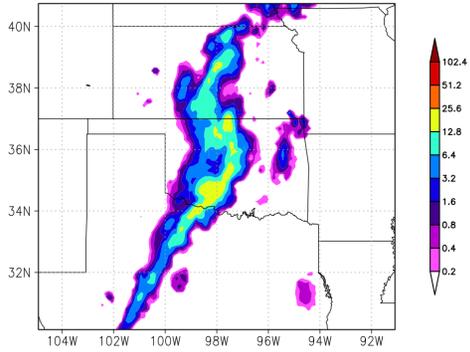
Ice, snow graupel and rain use inverse exponential distributions; cloud droplet uses gamma distribution.

# NLDAS-2

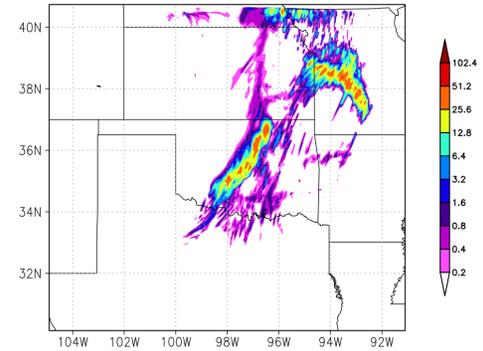
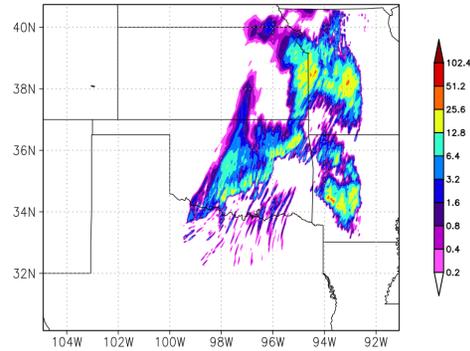
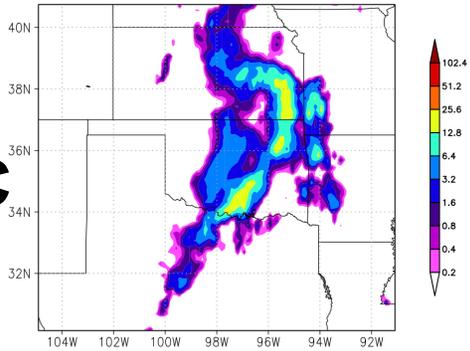
# WRF(Morrison)

# WRF(SBM)

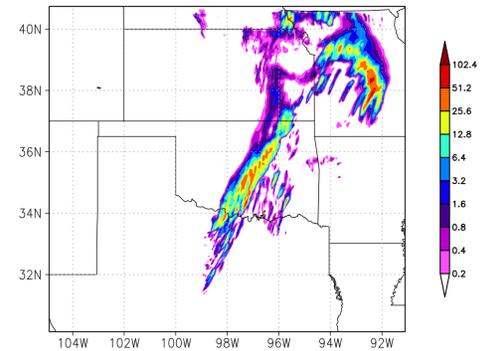
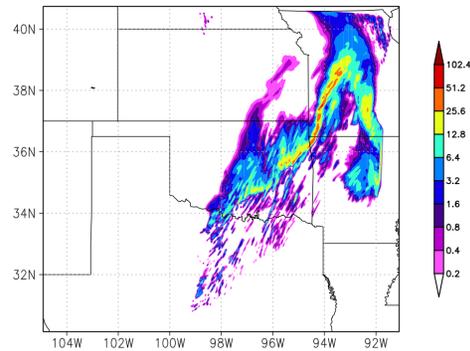
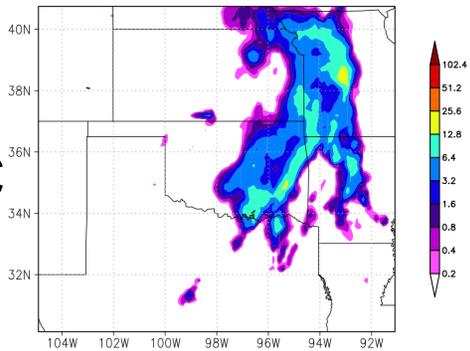
11UTC



14UTC



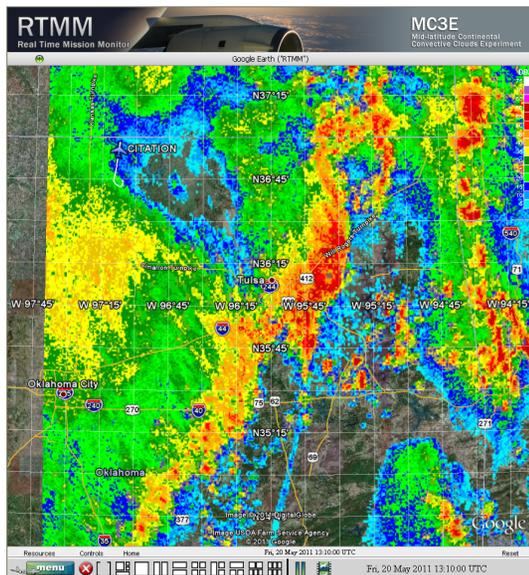
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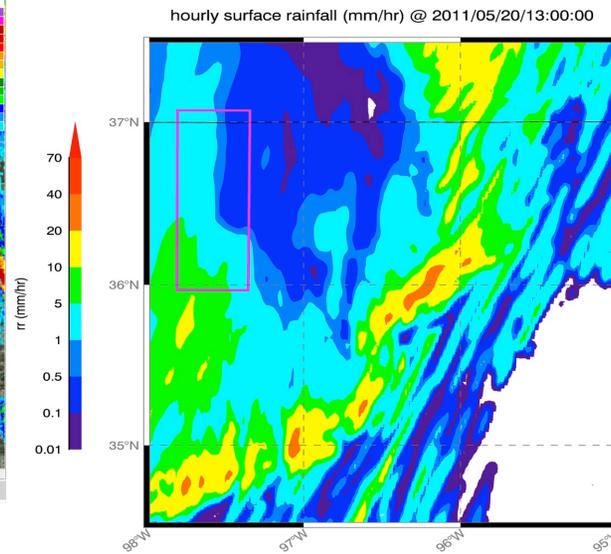
## Surface Rainfall Comparisons

# Comparison with In-situ PSD data

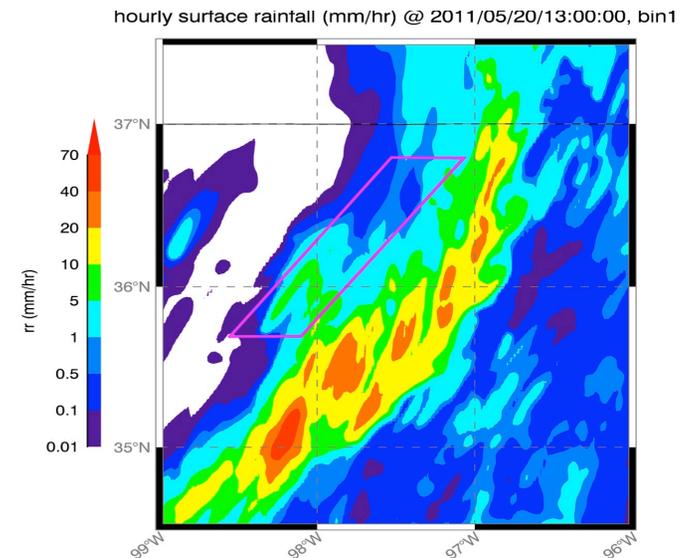
## Citation flight track



## Morrison scheme sample



## SBM scheme sample



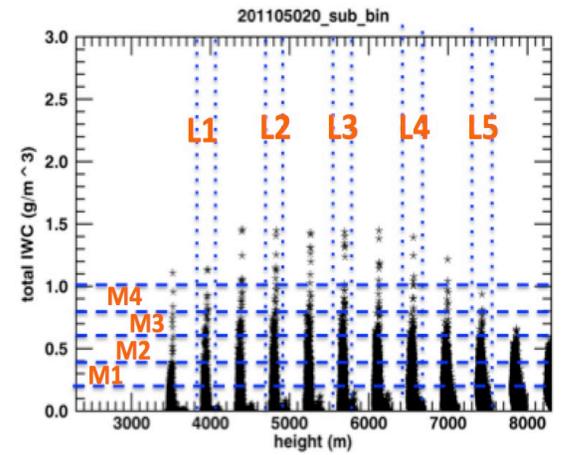
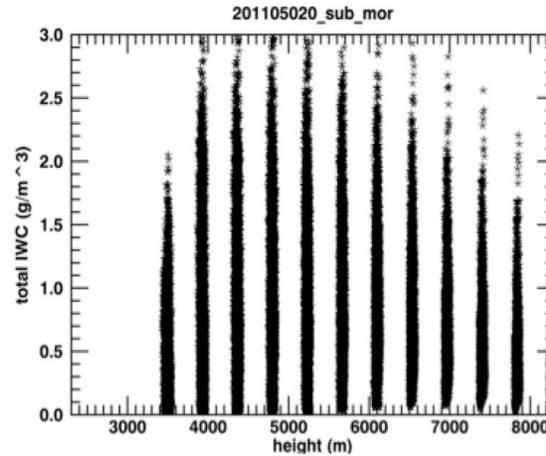
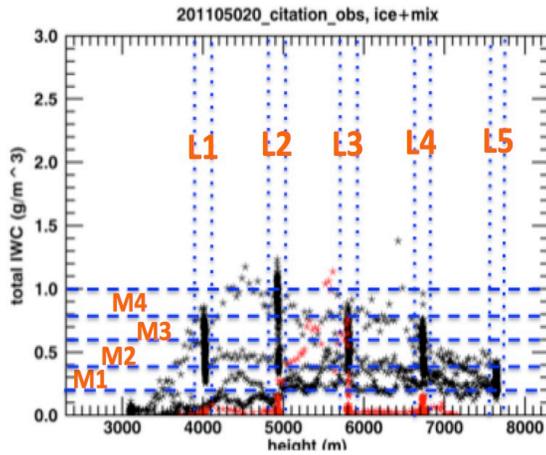
# Comparison with In-situ PSD data

## Citation Obs

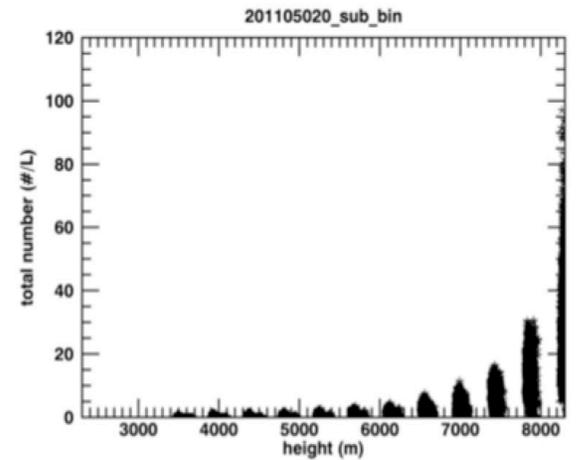
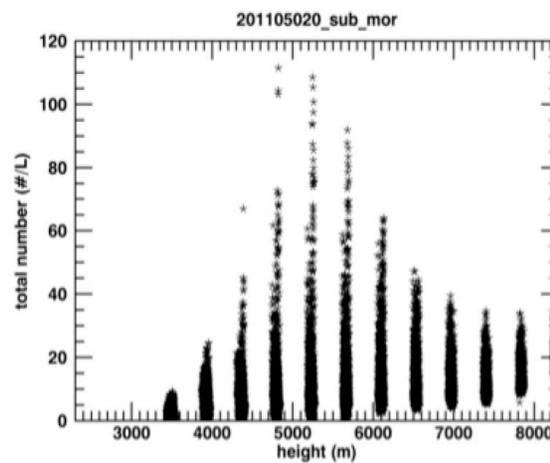
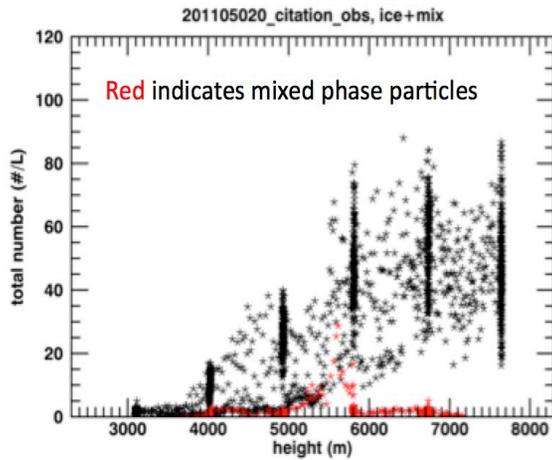
## WRF Morrison Sim

## WRF SBM Sim

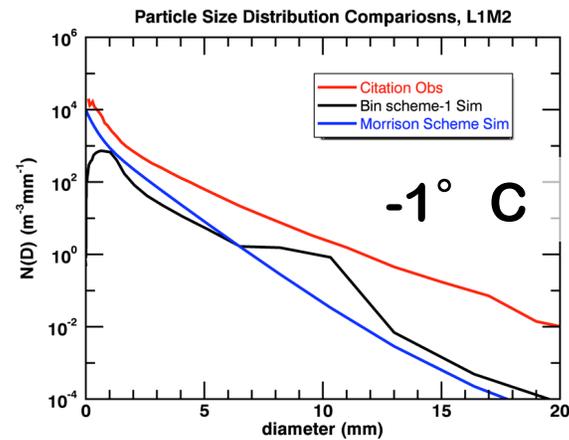
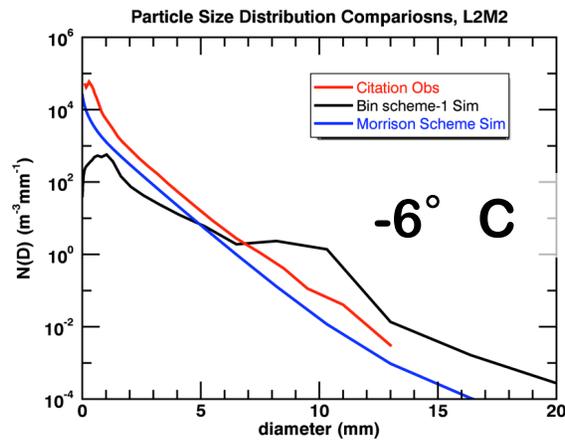
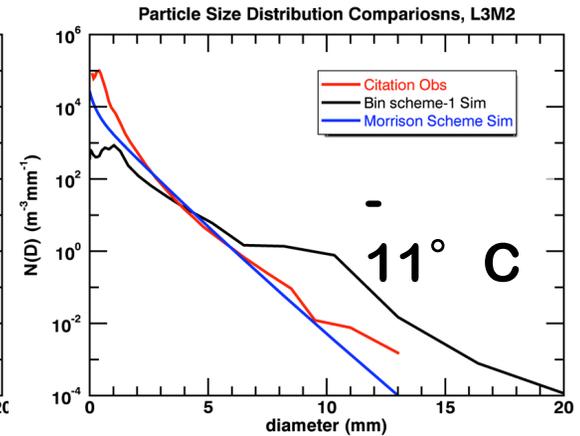
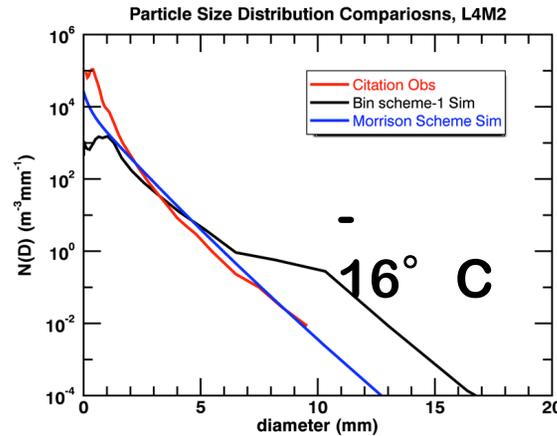
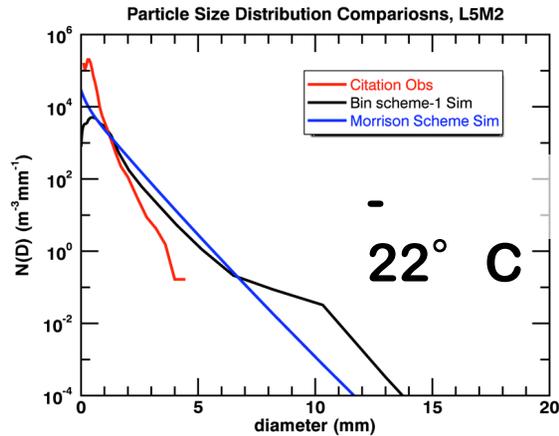
IWC



Total Num



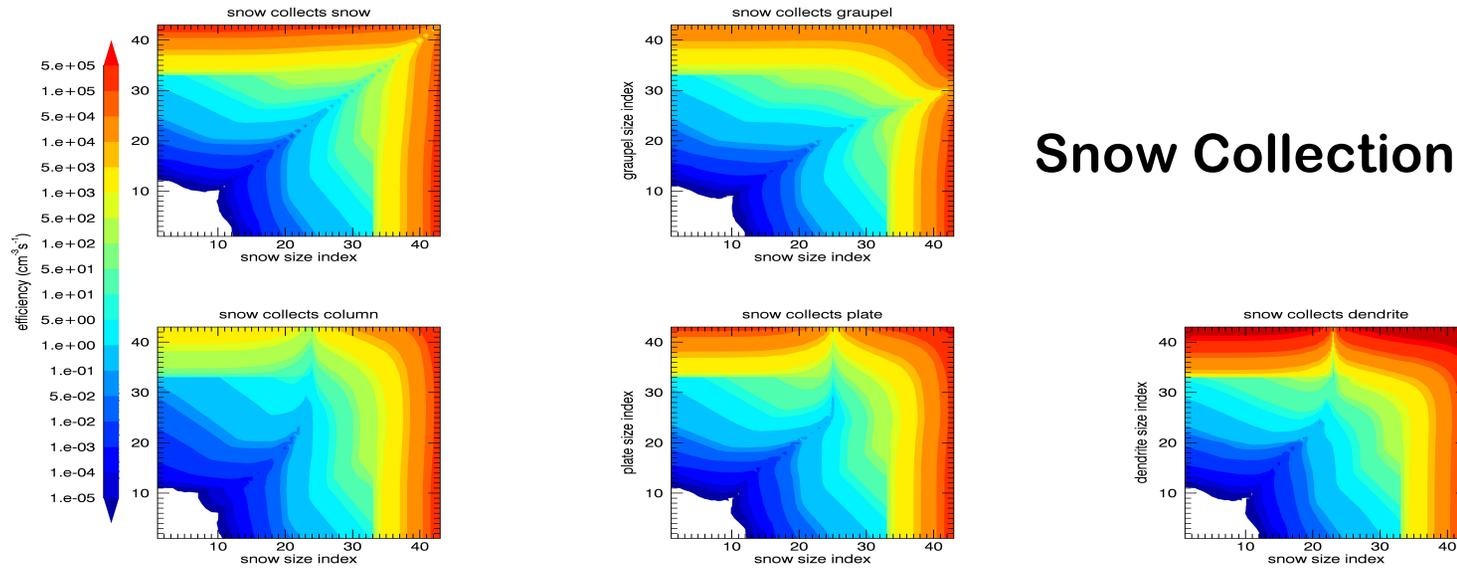
# Comparison with In-situ PSD data



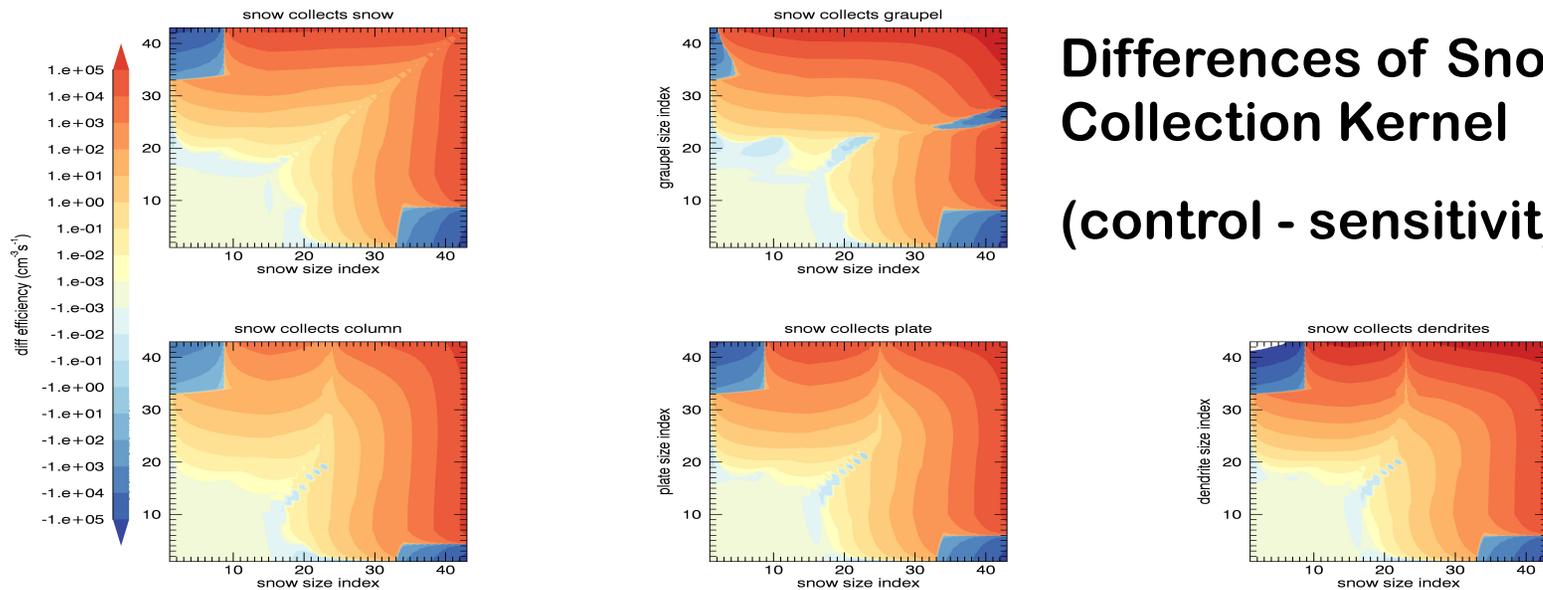
Comparisons indicate that both microphysical schemes have too strong collection.

For IWC between 0.4 ~ 0.6  $g/m^3$

# Sensitivity Test with Reduced Ec

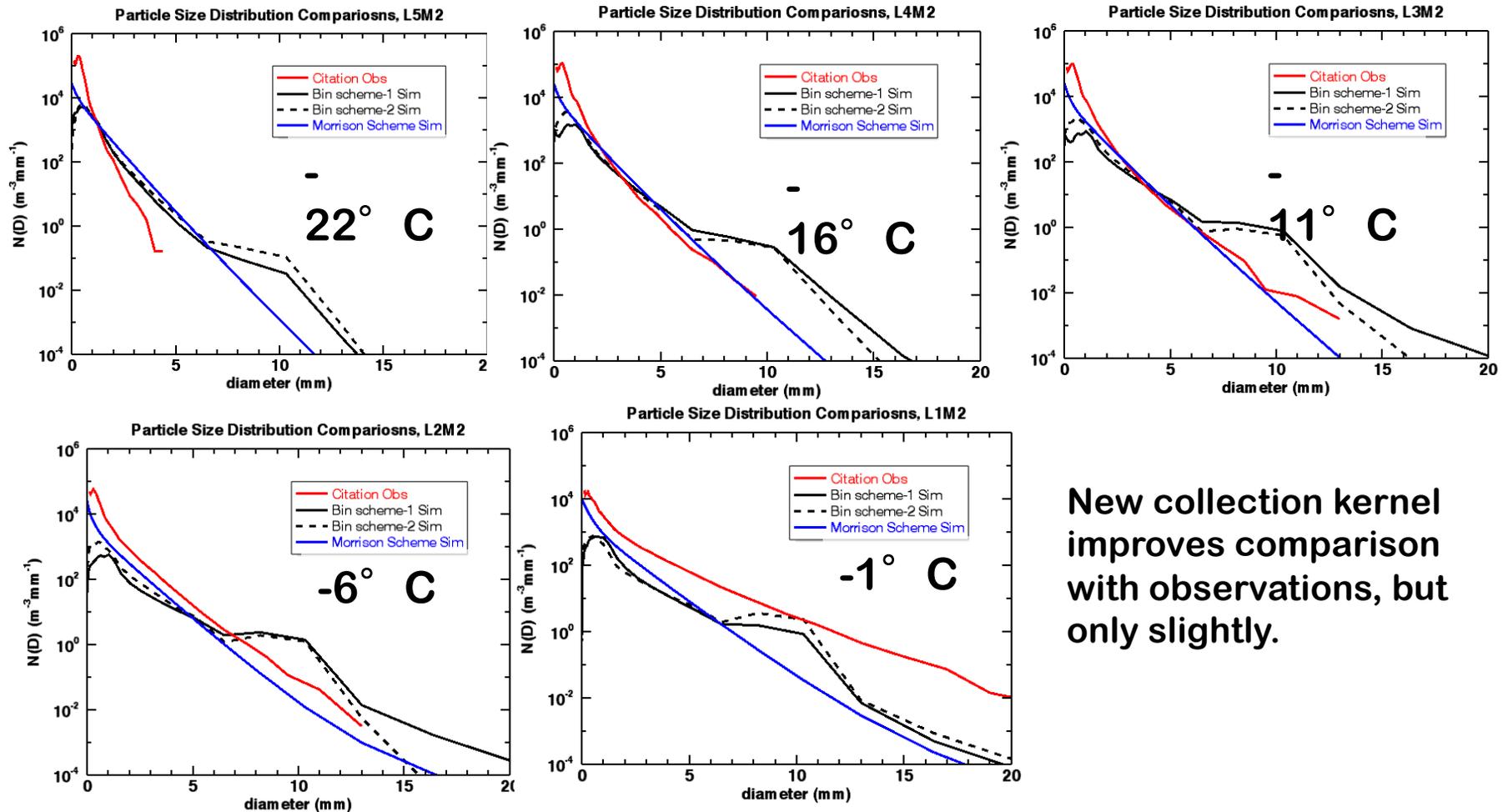


Snow Collection Kernel



Differences of Snow Collection Kernel  
(control - sensitivity)

# Sensitivity Test with Reduced Ec



New collection kernel improves comparison with observations, but only slightly.

For IWC between 0.4 ~ 0.6 g/m<sup>3</sup>

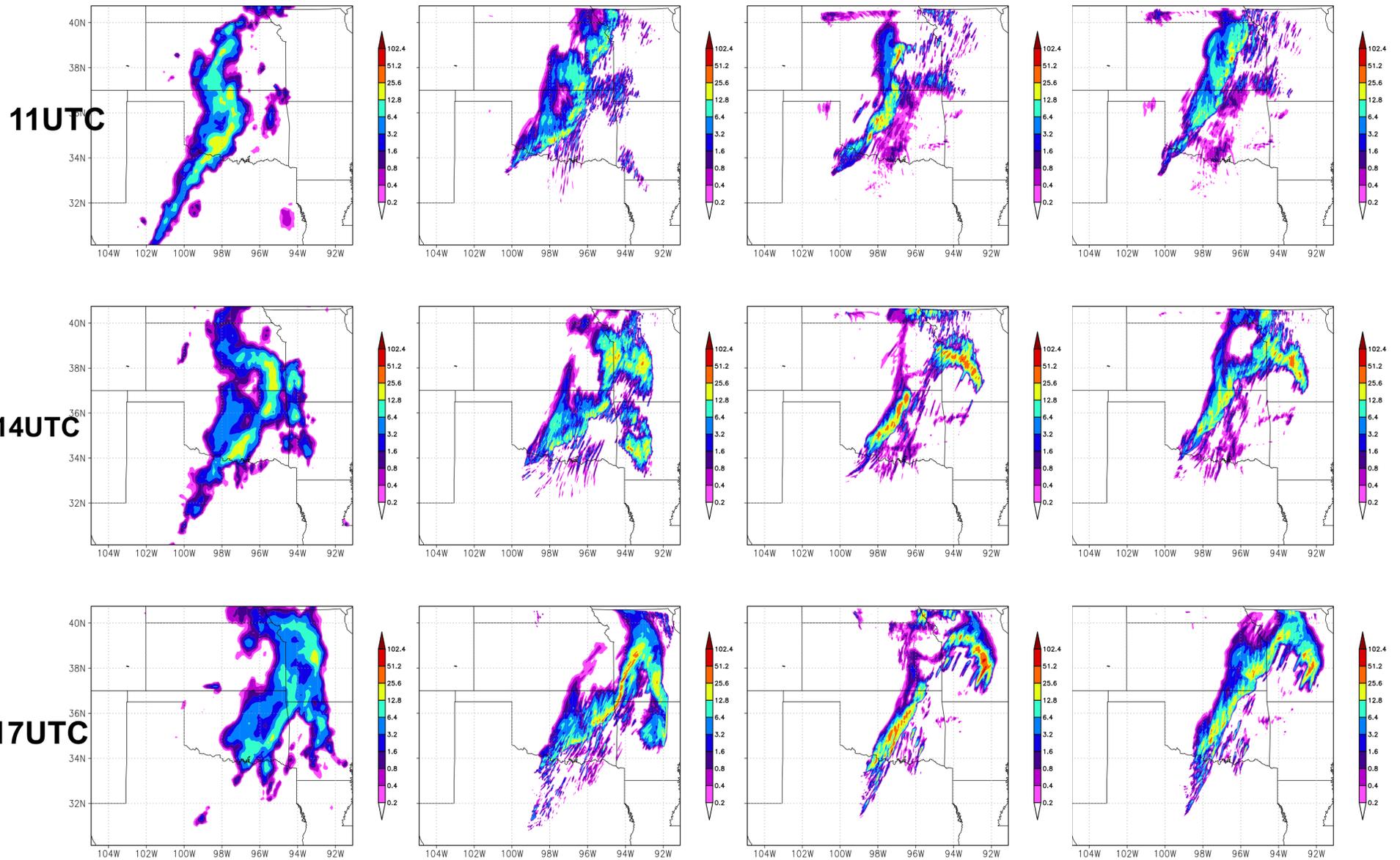
# Sensitivity Test with Reduced Ec

## NLDAS-2

## WRF(Morrison)

## WRF(SBM)

## WRF(SBM2)



# Reducing $E_c$ (T) produces better Radar CFADs Comparisons

C-Band

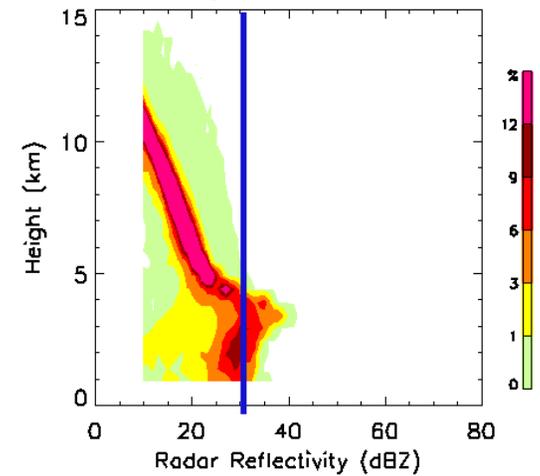
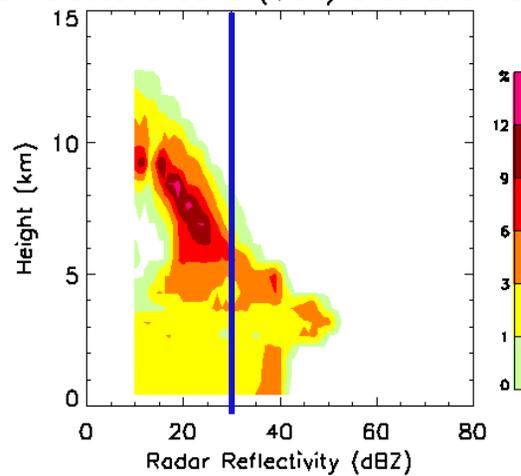
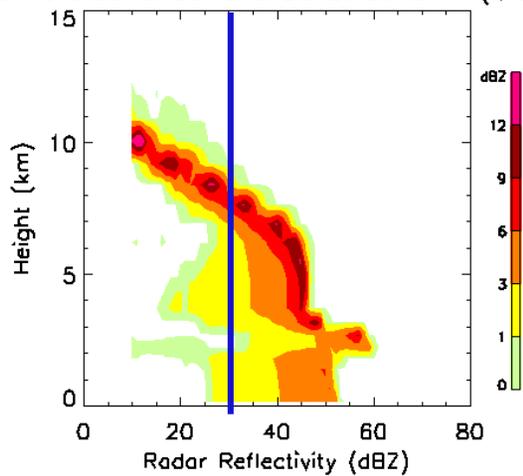
Original

Reduce  $E_c$

Observed

simulated C band cfad stratiform (MGA) simulated C band (MGA) cfad stratiform

PRESTORM cfad C BAND strat

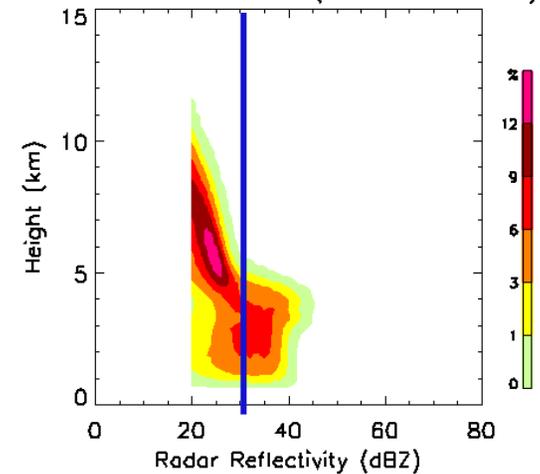
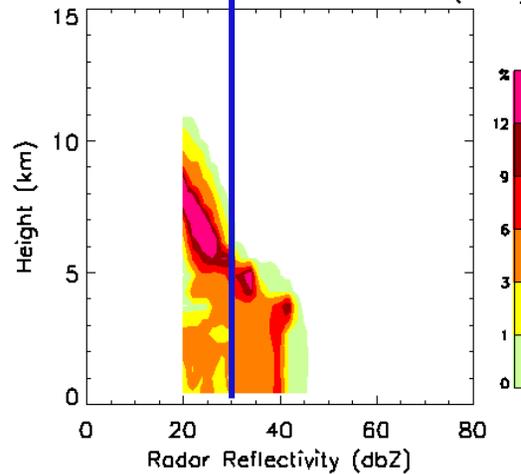
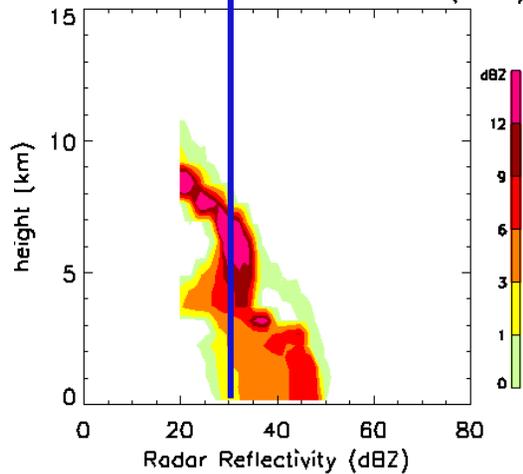


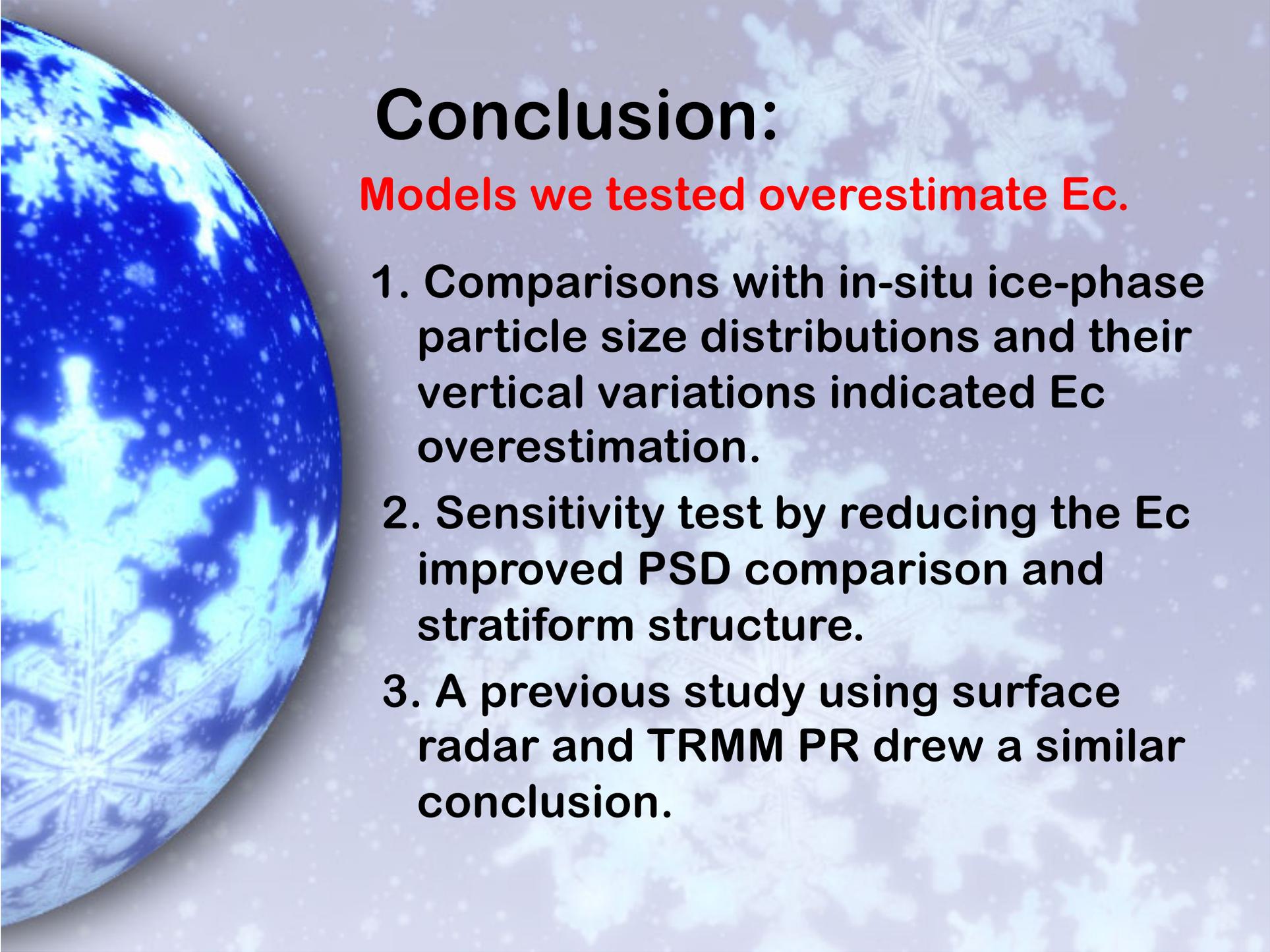
TRMM PR

simulated PR cfad stratiform (MGA)

simulated PR cfad stratiform (MGA)

PR cfad stratiform (total 18 cases)





# Conclusion:

**Models we tested overestimate  $E_c$ .**

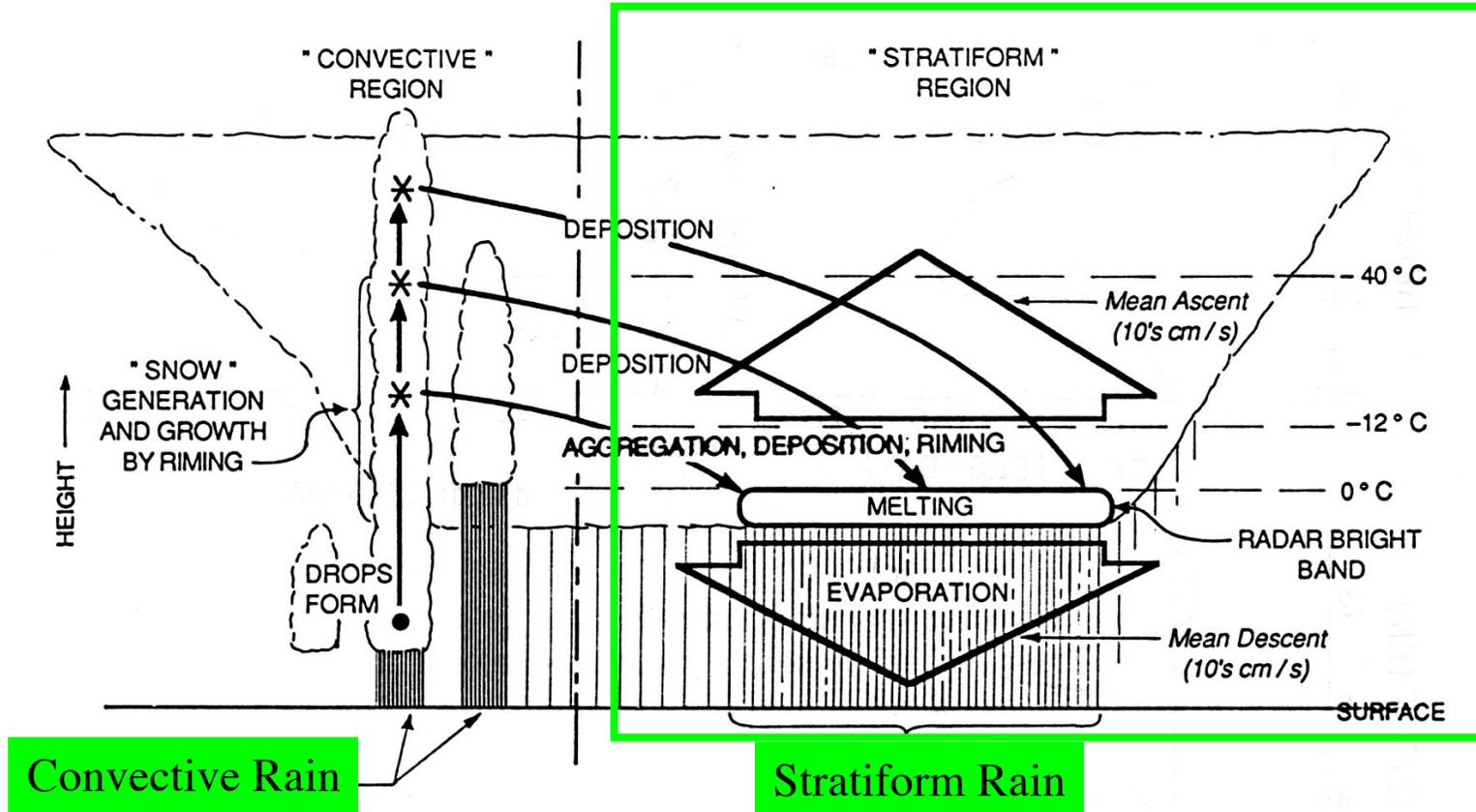
1. Comparisons with in-situ ice-phase particle size distributions and their vertical variations indicated  $E_c$  overestimation.
2. Sensitivity test by reducing the  $E_c$  improved PSD comparison and stratiform structure.
3. A previous study using surface radar and TRMM PR drew a similar conclusion.



# Future Plans:

1. Check rain DSD and see if the signals of large particles show.
2. Sensitivity tests by varying  $E_c(T)$  in the model.
3. Compare radar reflectivity profiles with airborne and GPM DPR.
4. Extend case studies to IPHEX and OLYMPEX.

# MCSs have extensive, homogeneous stratiform region



Houze, 1989: Observed structure of mesoscale convective systems and implications for large-scale heating. *Quart. J. Roy. Meteor. Soc.*, **115**, 425-461.