



# Does Snow Follow an $R - D_m$ Relationship?

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## Motivation

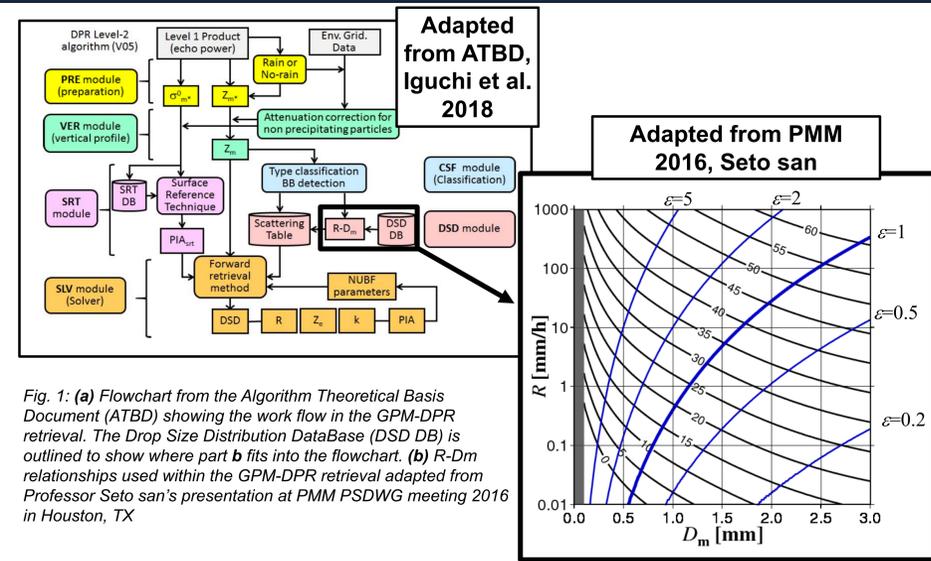
### GPM-DPR estimates of snowfall rate are low

Skofronick-Jackson et al. (2019) showed that even considering the sampling differences between CloudSat and GPM-DPR, the DPR global average snowfall rate was ~ 47% lower than those retrieved from CloudSat

Petersen et al. (2019) showed that when compared to surface based radar estimations in Finland, GPM-DPR had a 55% low bias.

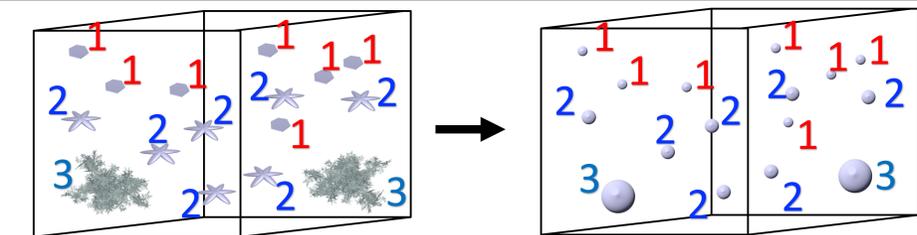
As a first investigation of where the low bias originates, we pose the question, "Does snow follow an  $R - D_m$  relation?"

## GPM-DPR algorithm V06



- Since the release of V04 of the GPM-DPR data, the DSD module has changed to use an  $R - D_m$  framework for retrievals
- The drop size distribution database (DSD DB), from which the  $R$  and  $D_m$  are retrieved, was created based on surface observations of rain
- This framework is used throughout the atmospheric column regardless of hydrometeor phase (see page 69, section 3.6.2-8 in ATBD)

## Derivation of Liquid Equivalent $D_m$



$$D_m = \frac{\int m(D)D N(D)dD}{\int m(D)N(D)dD}$$

$$D_{m,melted} = \frac{\int m(D)D_{melted} N(D)dD}{\int m(D)N(D)dD}$$

$$D_m = \frac{\sum_{i=0}^{n_{bins}} m(D_i)D_i N(D_i)dD_i}{\sum_{i=0}^{n_{bins}} m(D_i)N(D_i)dD_i}$$

$$D_{m,melted} = \frac{\sum_{i=0}^{n_{bins}} m(D_i)D_{melted,i} N(D_i)dD_i}{\sum_{i=0}^{n_{bins}} m(D_i)N(D_i)dD_i}$$

- In order to compare the  $R - D_m$  relations in the ATBD and those derived in snow, a liquid equivalent  $D_m$  needs to be calculated.
- We assume all particles in each bin are of the same density/type, then  $N(D_i)dD_i = N(D_{melted,i})dD_{melted,i}$ . This preserves mass content and total number concentration

## BAECC 2014 – 2015

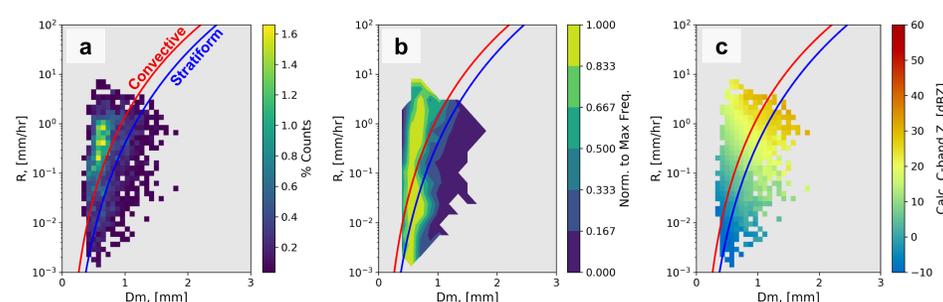


Fig. 3: Surface PIP snow observations (5-min) from Finland framed in the  $R - D_m$  space.  $R$  and  $D_m$  are in liquid equivalent. The solid blue and red lines correspond to the default stratiform and convective relationships in the GPM-DPR algorithm if  $\epsilon = 1$  (a) 2D histogram normalized to the total number of observations ( $n = 3020$ ) (b) Contoured 2D histogram (same as a) but normalized to the mode in each vertical bin as opposed to the total number of observations (c) Same 2D histogram as a, but now colored by the bin median C-band reflectivity. C-band reflectivity is calculated using the same expression as in von Lerber et al. (2017).

- Surface measurements of snowfall from southern Finland (61.485°N, 24.287°E) were collected by a Precipitation Imaging Package (PIP) between February 2014 – April 2015
- Bin-resolved mass for each PSD was retrieved following von Lerber et al. (2017) technique which allowed for the calculation of the liquid equivalent  $D_m$  and  $R$ .
- Figure 3 shows that snow PSDs over southern Finland tend to deviate from the prescribed relations, but still have increasing  $Z$  as  $R$  and  $D_m$  increase concurrently.

## Marquette, MI 2015 – 2018

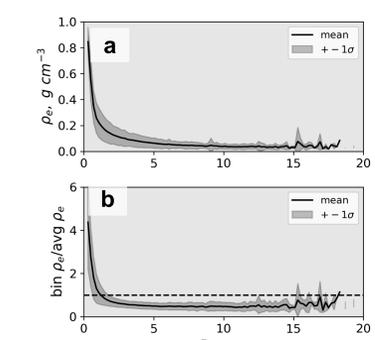
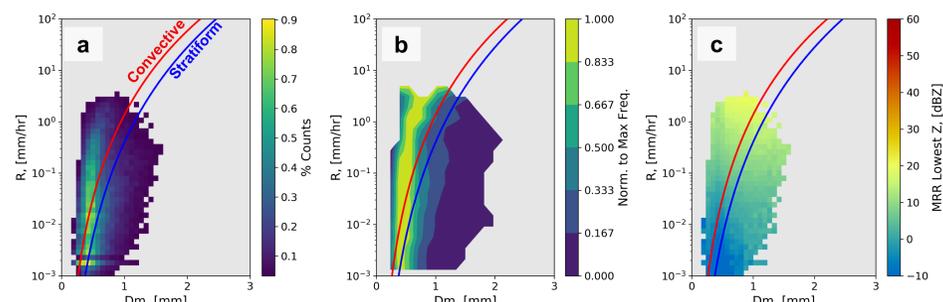


Fig. 4: Effective density ( $\rho_e$ ) adjustment. (a) Bin-wise  $\rho_e$  calculated from the retrieved masses from the PIP data in BAECC 2014-2015. (b) Ratio of the bin  $\rho_e$  to the total PSD average  $\rho_e$ .

- Surface measurements of snowfall in Marquette, Michigan (46.53°N, 87.55°W) have been taken every winter since 2014 using the PIP
- The bin-resolved mass is not readily available for this dataset, but an estimate of the 1-min PSD average effective density ( $\rho_e$ ) is provided for data extending back to 2015.
- The  $\rho_e$  is estimated using an empirical model relating the particle size and terminal fall speed (personal communication: Claire Pettersen)
- Using the relationship of PSD average  $\rho_e$  and that of the bin-wise  $\rho_e$  from the BAECC 2014-2015 data, a bin-resolved mass can be diagnosed for all PSDs collected (Fig. 4)



- After diagnosing the bin resolved mass as explained above, the Marquette data can be framed similarly to the BAECC observations
- Although the confidence of bin resolved mass is lower for the Marquette data compared to the BAECC data, Fig 5a,b shows that the Marquette data also deviates from the prescribed  $R - D_m$  relation
- Collocated MRR data demonstrate a consistent increase with  $Z$  away from the origin

## Comparison to Rain

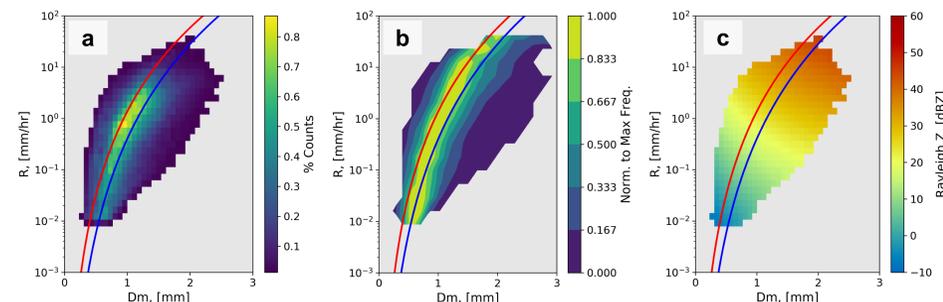


Fig. 6: As in Fig. 3, but now using 1-min rain observations using the 2-dimensional video dendrometers (2DVD) located at NASA GV campaigns in chronological order; LPVEX (Southern Finland), MC3E (Oklahoma, USA), IFloodS (Iowa, USA), IPHEX (North Carolina, USA), OLYMPEX (Coastal Washington, USA) and Wallops (Coastal Maryland, USA).  $n = 243,026$

- Fig. 6 shows that NASA GV 2DVD rain data are consistent with the prescribed relations, shifting from the stratiform curve (blue) to the convective curve (red) as  $R$  increases

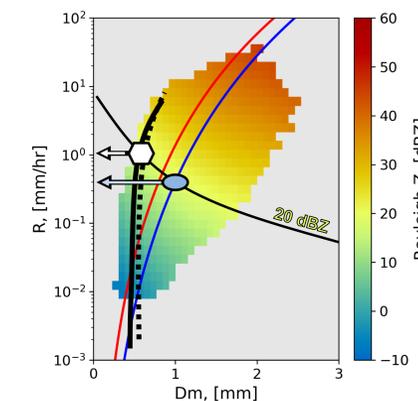


Fig. 7: Considering an example retrieval for a radar gate with a 20 dBZ echo. Shaded grid is the same as Fig. 6c. Thick black lines are new relations drawn from Fig. 3b (dotted) and Fig. 5b (dashed). Hexagon is the result using the new relations, oval is the old relations.

- For a given reflectivity, say 20 dBZ, if the prescribed relations are used, it would result in 0.4 mm hr<sup>-1</sup>, while if a new prescribed relation based off of the snow PSDs would result in 1 mm hr<sup>-1</sup>.
- This difference is ~50% bias, which is consistent with the literature reported estimates for how low GPM-DPR estimates are
- Better agreement is found with smaller  $R$  and  $D_m$

## Conclusions

- Ground based snow PSDs suggest that the  $R - D_m$  for rain does not apply well for snow.
- Using the rain relation as opposed to the snow relation would result in a 50% error for a radar echo of 20 dBZ.

## References

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