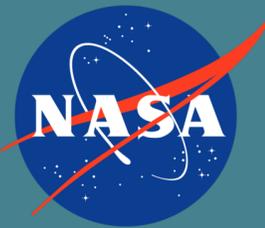


Diurnal variation of rainfall and the meridional overturning over West Africa during the pre-monsoon season

Africa during the pre-monsoon season

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1. Introduction

- West African precipitation mainly depends on the moisture transported inland from the Gulf of Guinea at seasonal and daily scales.
- During the pre-monsoon season, from April to June (AMJ), deep cloud systems with strong precipitation are predominant offshore in the Gulf of Guinea at around 5°N and are accompanied by a deep meridional circulation. Inland, at 15°N, a dry shallow meridional circulation is observed (Fig. 1).
- Two main features play important roles in the diurnal cycle of the West African vertical motion and precipitation climatology: the nocturnal low-level jet (NLLJ, Parker et al. 2005) and the sea-breeze front (Grams et al. 2010).

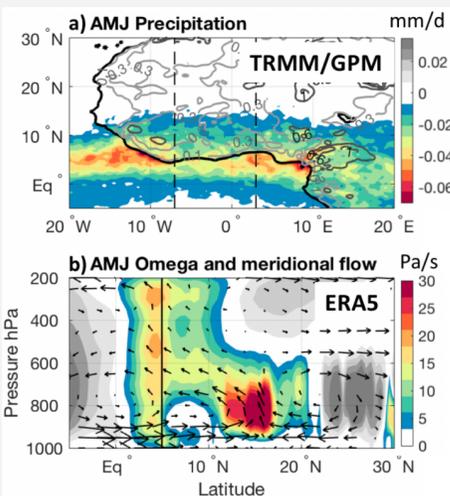


Fig. 1 a) PR-DPR precipitation during AMJ for 1998-2019, and b) ERA5 latitudinal cross section from 7°W-3°E (i.e., between vertical dashed lines in a) for omega and meridional flow.

2. Data

- Horizontal and vertical wind fields from ERA5 and MERRA-2 reanalyses interpolated to 0.5° resolution
- TRMM and GPM satellite radar precipitation gridded at 0.25° and categorized as (deep) convective, stratiform, and shallow (convective) from 2APR and 2ADPR products
- All data shown is during AMJ (the pre-monsoon) for 1998-2019

3. Results

3.1 Diurnal cycle of vertical velocity from ERA5

Diurnal Features:

- a) **Sea Breeze**
- b) **Nocturnal low-level jet (NLLJ)**
- c) **Mixing Layer**

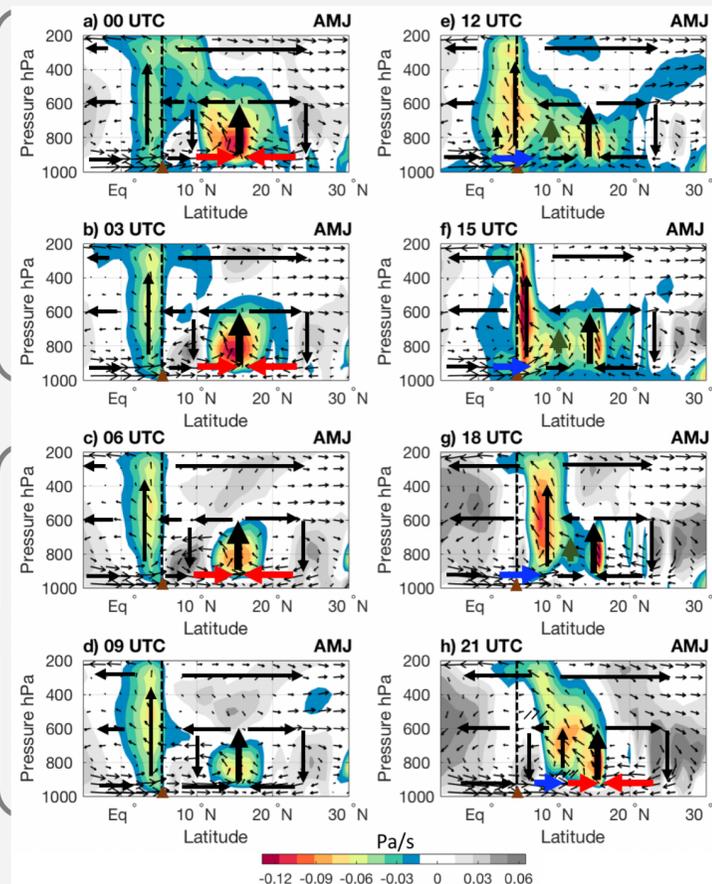
00 – 03 UTC

- NLLJ very active, enhances Saharan shallow meridional circulation
- Deep upward motion over ocean

06 – 09 UTC

- NLLJ and shallow meridional circulation decaying
- Deep upward motion over ocean strengthening

Fig. 2 Diurnal variation of vertical motion (Pa/s) and meridional flow in latitudinal cross sections from 7°W-3°E during AMJ. The solid line at 5.5°N highlights the coastline.



12 – 15 UTC

- Deep upward motion propagates onto land due to sea-breeze front (11.5 m/s)
- Well-mixed air in planetary boundary layer
- Shallow meridional circulation is weak

18 – 21 UTC

- Deep upward motion propagates inland (10°N) due to sea-breeze front
- Deep convection interacts with the southern part of dry Saharan shallow meridional circulation

3.2 TRMM-GPM rain types

Deep convective and stratiform rain propagates inland from 12 to 18 UTC due to sea-breeze front. Shallow convective rain stays closer to coast.

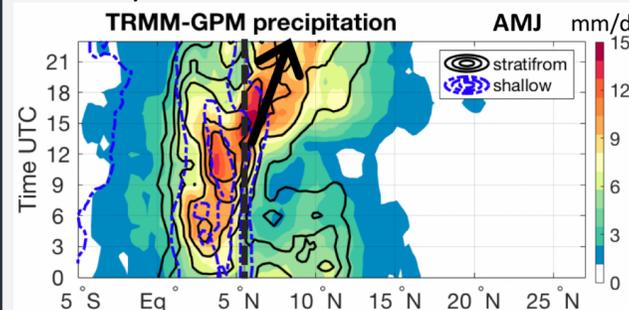


Fig. 3 Convective (shaded), stratiform (black contours each 1 mm/d starting at 2 mm/d), and shallow (blue contours each 0.5 mm/d starting at 0.5 mm/d) rain in mm/d from PR-DPR during AMJ.

3.3 MERRA-2 vs ERA5

- MERRA-2 doesn't show inland propagation of deep convection (represented by 400-hPa omega) between 12 to 21 UTC as suggested by ERA5
- MERRA-2 shows sea breeze (represented by 850-hPa omega) with slower phase speed than ERA5 and is not linked to 400-hPa omega and deep convection

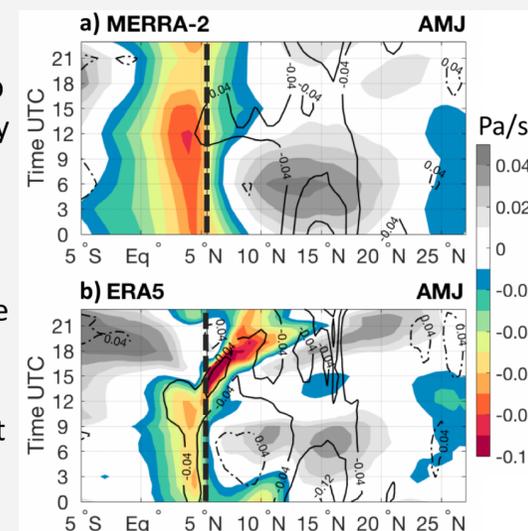


Fig. 4 Latitude-time diagrams of vertical velocity at 400 hPa (shading) and 850 hPa (contours) in Pa/s from a) MERRA-2 and b) ERA5 during AMJ.

4. Conclusions

- The meridional overturning circulation over West Africa during AMJ has strong diurnal cycle linked to the sea breeze and NLLJ.
- The sea breeze drives deep convection from the coast into land during the early afternoon.
- The continental NLLJ produces large low level-convergence and strengthens the dry shallow cell at night.
- MERRA-2 overestimates convection over coastal region and does not show inland propagation of deep convection, while ERA5 captures this well.

5. Acknowledgments

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6. References

- Parker et al. 2005. The diurnal cycle of the West African monsoon circulation. *Quarterly Journal of the Royal Meteorological Society*, 131(611), 2839-2860.
- Grams et al. 2010. The Atlantic inflow to the Saharan heat low: Observations and modelling. *Quarterly Journal of the Royal Meteorological Society*, 136(S1), 125-140.